



A Late Permian paleopole from the Ikakern Formation (Argana basin, Morocco) and the configuration of Pangea

Dennis V. Kent^{a,b,*}, Paul E. Olsen^a, Giovanni Muttoni^c, Mohammed Et-Touhami^d

^a Lamont-Doherty Earth Observatory of Columbia University, Palisades, NY 10964, USA

^b Earth and Planetary Sciences, Rutgers University, Piscataway, NJ 08854, USA

^c Dipartimento di Scienze della Terra 'Ardito Desio', Università degli Studi di Milano, via Mangiagalli 34, I-20133 Milan, Italy

^d 2GPMH, Département des Sciences de la Terre, Université Mohammed Premier Oujda, 60000 Oujda, Morocco

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ABSTRACT

The nearly 2000 m-thick Ikakern Formation is the basal sedimentary unit in the Argana basin of Morocco, consisting of fluvial red and purple conglomerates in the lower part and interbedded red conglomerates, sandstones, siltstones and mudstones in the upper part. It unconformably overlies deformed and metamorphosed Variscan basement rocks and is unconformably overlain by the finer grained lacustrine to fluvial sediments of the Early to Late Triassic Timesgadiouine Formation. Vertebrate fossil and chronostratigraphic constraints indicate a post-Kiaman, Late Permian age for at least the sampled upper member (t2) of the Ikakern Formation. A stable, high unblocking temperature dual polarity characteristic magnetization that passes a local fold test gives a paleomagnetic pole for 14 sites at $48.8^{\circ}\text{N } 246.3^{\circ}\text{E } A95 = 7.2^{\circ}$ after correction for inclination flattening and indicates a paleolatitude of 14°N for the nominal sampling location at $30.75^{\circ}\text{N } 9.10^{\circ}\text{W}$. The Ikakern pole agrees well with poles of similar age deemed reliable in a recent assessment of published data from South America, Africa, Adria, and Australia, which together ($N = 7$) provide a mean Gondwana pole for an age bin centered on 260 Ma at $52.2^{\circ}\text{N } 239.8^{\circ}\text{E } A95 = 5.0^{\circ}$ (NW Africa coordinates). In conjunction with a robust mean 260 Ma pole from Laurasia, these results confirm a Pangea A configuration for the Late Permian. This is in contrast to some recently published interpretations of paleomagnetic data from the Gondwanides of South America that would prolong the existence of a Pangea B configuration from the Late Carboniferous–Early Permian through the Late Permian but which instead probably reflect effects of remagnetization or local rotation.

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1. Introduction

For about the last 20 million years (Myr) of the Paleozoic, a time frame encompassing the Middle and Late Permian Epochs (272.3–252.2 million years ago or mega-annum (Ma); Gradstein et al., 2012), a recent assessment of global paleomagnetic poles restricted to igneous or flattening-corrected sedimentary data from cratonic regions shows there are only six reference poles deemed reliable for an assembled Gondwana and ten entries for Laurasia (North America and Europe) (Kent and Muttoni, 2020). These Gondwana and Laurasia results give mean poles for 260 Ma that are compatible with a Pangea A-type configuration for the Late Permian, in agreement with most recent studies (e.g., Domeier et al., 2011a) even though the Pangea B configuration of Irving (1977) had sometimes been thought to extend into the Triassic (Morel and Irving, 1981; Torcq et al., 1997). However, some confusion has emerged for Middle to Late Permian paleomagnetic results from South America, which Domeier et al. (2011b) argue

support a Pangea A-type configuration whereas Tomezzoli et al. (2018) more recently advocate a Pangea B configuration persisting through the Late Permian. The small number of reliable entries and the modest overall precision ($A95 = 5.7^{\circ}$) of the Gondwana mean pole for 260 Ma (Kent and Muttoni, 2020) make it desirable to obtain additional results especially from NW Africa, which should have faced eastern North America in a conventional Pangea A reconstruction but for which no reliable Late Permian poles are available.

Accordingly, we report new paleomagnetic results from the Ikakern Formation, a sedimentary red bed unit of Late Permian age exposed in the Argana Basin of western Morocco in NW Africa. The characteristic magnetization (ChRM) isolated by thermal demagnetization passes a local fold test, indicating acquisition prior to the unconformably overlying Timesgadiouine Formation of Early to Late Triassic age, and have normal and reverse polarities, which in conjunction with vertebrate biostratigraphy of Permian aspect (Jalil et al., 2002), constrain the magnetization age to the latest Kiaman to Illawara polarity interval range of the Late Permian, essentially 265–252 Ma. The Ikakern paleomagnetic pole corrected for inclination flattening can then be compared to compiled

* Corresponding author.

E-mail address: dvk@ldeo.columbia.edu (D.V. Kent).

reference poles for the interval 270–250 Ma from various entities that comprised Gondwana, especially South America, and with the mean 260 Ma pole for Laurasia to test the basic Pangea configuration for the Late Permian.

2. Geologic setting and sampling of the Ikakern Formation

Following the regional mapping and lithostratigraphic column of (Tixeront, 1973), the continental strata of the Argana rift basin in western Morocco (Fig. 1) can be divided into four tectonostratigraphic sequences (Olsen, 1997; Olsen et al., 2000) (Fig. 2). The Ikakern Formation (t1 and t2 of Tixeront (1973)) constitutes the oldest tectonostratigraphic sequence (TS-I) and is comprised of a 1000+ meter-thick basal conglomerate (t1) that is restricted to a graben structure in the center of the Argana Valley and followed by several hundred meters of interbedded conglomeratic sandstones, siltsones and mudstones (t2) that are generally reddish in colour and include mudcracks and other evidence of exposure (Brown, 1980). The Ikakern Formation is overlain with a locally marked angular unconformity by the

Timesgadiouine Formation (t3 to t5 of Tixeront, 1973) of tectonostratigraphic sequence TS-II, comprised of a thin basal conglomerate (t3) and a ~1000 m-thick sequence of reddish mudstones interbedded with red siltstones and sandstones that are sometimes coarse-grained (Brown, 1980). The Bigoudine Formation (t6 to t8 of Tixeront, 1973), which makes up TS-III and the base of TS-IV, conformably overlies the Timesgadiouine Formation and is comprised of a basal conglomerate and fluvial sandstone (t6) succeeded by laterally-persistent cycles of red mudstone and sand-patch fabrics with some black shales in the upper parts. Deposition of the ~1000 m-thick Bigoudine Formation ended just after the end-Triassic boundary with outpourings of the Argana Basalt in the latest Rhaetian. See also El Hachimi et al. (2011) and Leleu et al. (2016) for further information on the geology of the study area.

The age of the Ikakern is constrained by four lines of stratigraphic evidence: 1) ages of the underlying units; 2) age of an intruding body of igneous rock; 3) age inferred on the basis of the fossil vertebrates from the upper part of the formation itself; and 4) age of the unconformably overlying conglomerates.

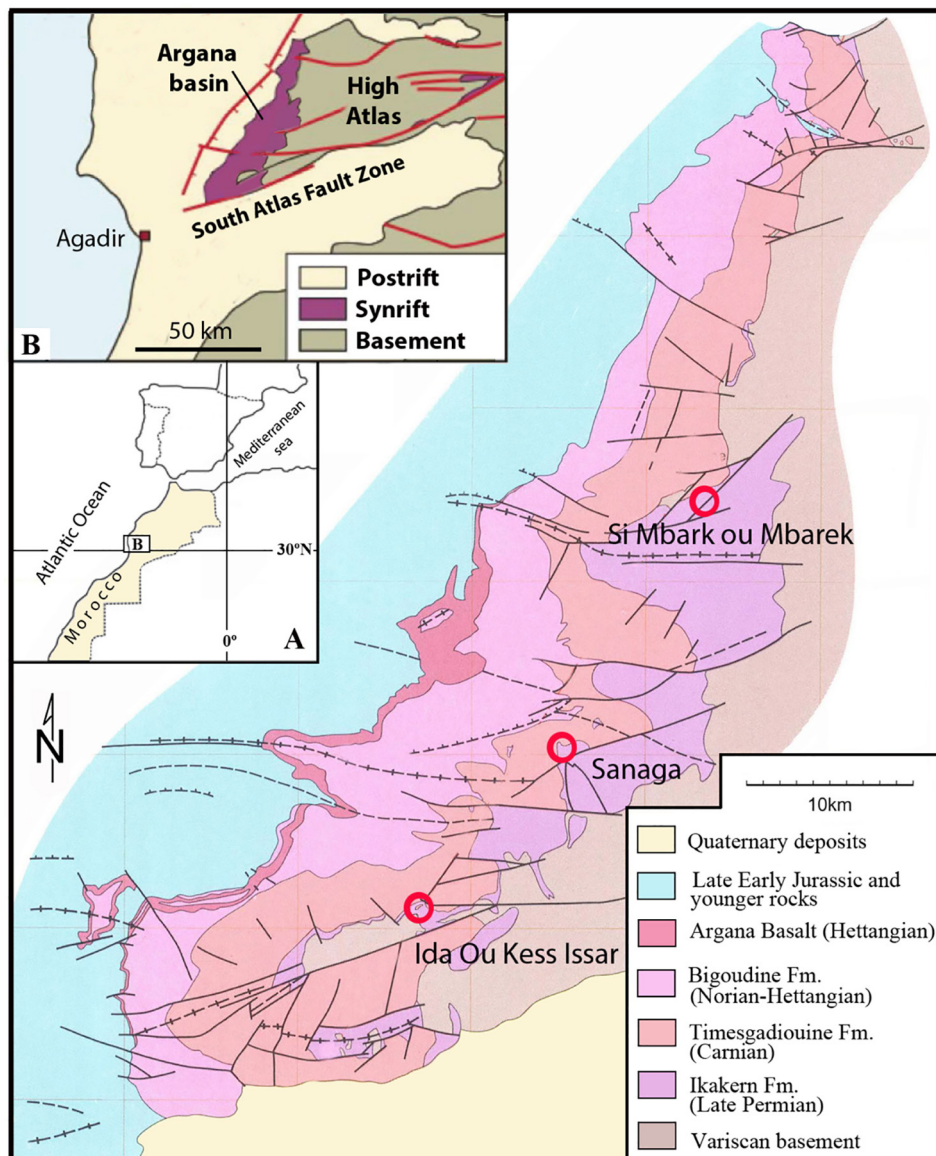


Fig. 1. Structural map of Argana Basin (modified after Tixeront, 1973) and paleomagnetic sampling localities in Ikakern Formation (see Table 2 for precise site coordinates and Fig. 2 for stratigraphic ranges).

Argana Basin Stratigraphy

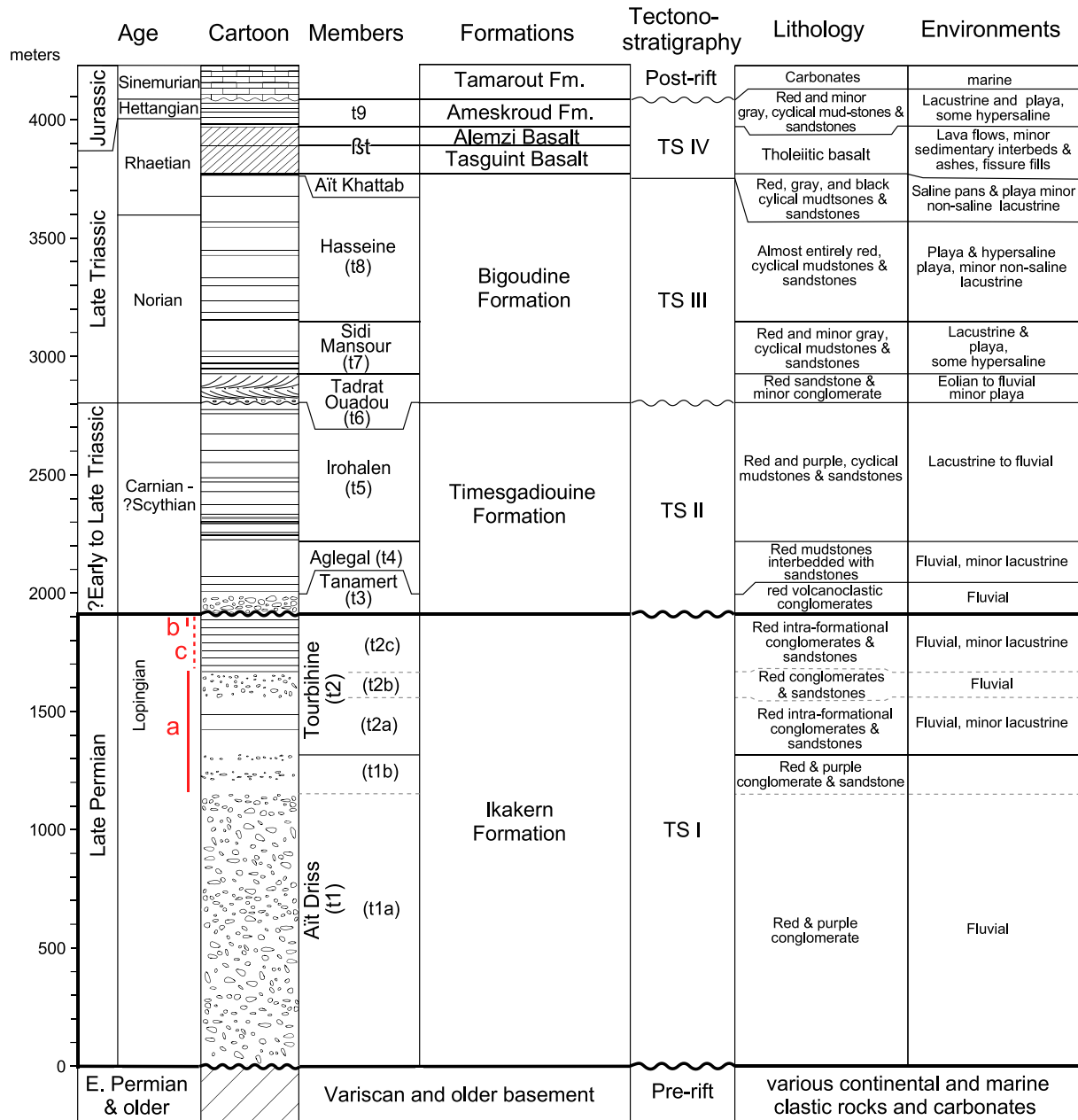


Fig. 2. Generalized stratigraphic section of Argana Basin (based on Brown, 1980; Deenen et al., 2010; El Hachimi et al., 2011; Olsen, 1997; Tixeront, 1973). Thicknesses for Bigoudine and Ikakern formations are based on our own measurements. Distribution of paleomagnetic samples are shown by red lines: a, Si Mbark ou Mbarek section; b, Sanaga section; c, Ida Ou Kess Issar section (dashed red line). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

1) The youngest rocks underlying the Ikakern are Late Carboniferous (? Early Permian) in age and are part of the pre-rift basement that was deformed and intruded during the Variscan Orogeny. Paleozoic and older clasts of quartzites, phyllites, limestone and less common vein-quartz, igneous, and sedimentary terrigenous rocks from the basement terrane comprise the conglomerates of the Ikakern (Brown, 1980). The locally basal Ikakern at our Ida Ou Kess Issar locality unconformably overlies continental sedimentary strata of the Oued Issène Formation, which in the Ida ou Ziki subbasin (Souss Basin) are Stephanian A/B - Kasimovian (Late Carboniferous) based on the flora and microflora (Hmich et al., 2006) and which has also produced tetrapod footprint assemblages of Late Carboniferous aspect (Lagnaoui et al., 2017) and some tetrapod skeletal remains

(Werneburg et al., 2019). Elsewhere, the Oued Issène Formation is capped by conglomerates of the Tirkou Formation identified by some authors as Early Permian (e.g., Roscher and Schneider, 2005). The Ikakern unconformably overlies similar Late Carboniferous to? Early Permian strata in the adjacent Ida Ou Zal sub-basin (Hmich et al., 2006; Lagnaoui et al., 2017) and thus must be younger than Late Carboniferous to? Early Permian.

2) The Tourbihine Member (t2) of the Ikakern Formation is intruded by a basaltic igneous body, the Aguerd Ouaooudid “neck” (Medina and El-Amrani, 1989), 2 km west (30.727872°, -9.082686°) of our sites at Sanaga and which is now recognized as part of Central Atlantic Magmatic Province (CAMP; Marzoli et al., 1999) activity (El Hachimi et al., 2011). The Amelal Sill that invaded the upper

Bigoudine Formation of latest Triassic age is dated by zircon U–Pb CA-ID-TIMS geochronology at 201.564 ± 0.045 (Blackburn et al., 2013) and 201.569 ± 0.042 (Marzoli et al., 2019) and was a feeder to the adjacent Alemzi Formation CAMP flow (Blackburn et al., 2013). As all CAMP rocks dated globally are within 1 Myr of this date, the Ikakern must be older than about 202 Ma (very latest Triassic) given its position near the base of a stratigraphic section essentially capped by the Argana CAMP basalt flows.

- 3) Vertebrate fossils, both skeletal remains and footprints, suggest a Late Permian age for the Tourbihine Member (t2) of the Ikaren Formation, although all of the taxa present are endemic at the species level and some suggest discrepant ages. Dutuit (1988) described the unmistakable hammer-headed lepospondyl amphibian *Diplocaulus minimus* from the upper Tourbihine Member. This genus is otherwise known from Late Carboniferous and Early Permian deposits of North America. Additional taxa include captorhinid *Acrodonta irerhi* (Dutuit, 1976) and a moradisaurine captorhinid (Jalil, 1999; Jalil and Dutuit, 1996) that support a Middle Permian age. However, the discovery of “advanced” large and distinctive globally distributed pareiasaur parareptile *Arganaceras vacanti* in the Tourbihine Member suggests a Late Permian age (Jalil and Janvier, 2005) [late Tatarian, or latest Permian of the Soviet time scale: Lucas and Shen, 2018]. The footprint record of the Tourbihine Member is consistent with this age assessment (Hminna et al., 2012; Voigt et al., 2010) with assemblages placed within the *Paradoxichnium* assemblage of inferred Late Permian age (Voigt and Lucas, 2018).
- 4) Overlying the Ikakern with an angular unconformity is the conglomeratic Tanamert Member (t3) of the Timezgadiouine Formation (Brown, 1980; Tixeront, 1973). There are several vertebrate-bearing levels in the Timezgadiouine Formation (Klein et al., 2010) with the upper strata (Irohale Member, t5) producing a relatively rich assemblage of skeletal remains and footprints considered to be Carnian in age (Jalil, 1999; Kammerer et al., 2012; Lagnaoui et al., 2016; Lagnaoui et al., 2012; Lucas, 1998; Olsen et al., 2000). The underlying Aglegal Member (t4) is of decidedly Middle Triassic aspect having footprint assemblages that include the iconic middle Triassic ichnite *Chirotherium barthi* (Klein et al., 2011) as well as scrappy skeletal remains (Olsen et al., 2000) and ostracodes and charophytes (Medina et al., 2001). Most importantly, the basal member of the Timezgadiouine Formation, the conglomeratic Tanamert Member (t3), lying unconformably on the Ikakern, has produced a footprint assemblage of Early Triassic aspect (Klein et al., 2010). Igneous clasts from the Tanamert Member have yielded legacy $^{40}\text{Ar}/^{39}\text{Ar}$ ages of 247 ± 5 Ma, 280 ± 15 Ma, and 300 ± 20 Ma (Jones, 1975). The Tanamert Member (t3) must be younger than these apparently Carboniferous to Triassic clasts, keeping in mind both the low precision and low accuracy of the available radioisotopic dates.

In summary, the best available evidence shows that the Ikakern must be older than Early Triassic and younger than Late Carboniferous, with internal biostratigraphic evidence suggesting a Late Permian age for at least the Tourbihine Member (t2) of the Ikaren Formation.

The Ikakern Formation was sampled during two sampling campaigns, the first in the year 1999 near Si Mbark ou Mbarek, and the second in 2001 at Ida Ou Kess Issar and Sanaga, as part of a reconnaissance paleomagnetic survey of the Argana basin (Figs. 1, 2). Fine-grained red siltstones and red sandstones were sampled wherever possible. Block samples oriented with a magnetic compass and clinometer were taken near the town of Si Mbark ou Mbarek; three or more block samples representing separate beds were grouped into sites that were collectively distributed over 200–300 m of stratigraphic section. Oriented cores taken with a portable gasoline-powered drill and 2.5 cm-diameter diamond bit were taken in a section exposed along a river near Ida Ou Kess Issar, and across a small antiform near Sanaga. Bedding generally dipped to the northwest (14° – 36°) but it was also possible to sample southeast dipping beds for a local fold test at Sanaga. In total, 58

oriented samples were successfully recovered from 15 sites at the three distinct localities for laboratory analysis.

3. Paleomagnetic analysis

Magnetic remanence measurements and thermal demagnetization experiments on standard 2.5 cm-diameter plugs cut from the block samples or sliced from the cores were made using a 2G Enterprises (Mountain View, CA, USA) Model 760 3-axis DC-SQUID cryogenic magnetometer and custom-built ovens all housed in a magnetically shielded room at the Lamont-Doherty Paleomagnetism Lab. Isothermal remanent magnetization (IRM) acquisition and thermal demagnetization experiments following the orthogonal method of (Lowrie, 1990) (2.5 T and 0.2 T in this case) show that high coercivity hematite, with maximum unblocking temperatures of around 685 °C, is the predominant carrier of remanence in these red beds with a more thermally-distributed phase with an unblocking temperature concentration at around 550 °C that may reflect a finer grain size pigmentary hematite fraction (Irving and Opdyke, 1965) (Fig. 3). The natural remanent magnetization (NRM) of the 58 samples averages 6 mA/m and has directions broadly streaked between northwest-down and southeast-up (Fig. 4A). Progressive thermal demagnetization of samples in 10 or more steps reveals that the NRM is typically comprised of various combinations of two or three components of magnetization including a relatively minor low unblocking temperature component A, a northerly and down component B with intermediate unblocking temperatures, and what we regard as a final high unblocking temperature component C with either northwest-down or southeast-up directions (Fig. 5). The directions of the linear demagnetization trajectories chosen from Zijdeveld orthogonal plots (typically three steps from 100° to 300 °C for the A component, four or five steps from 300° to 600 °C for the B component, and three to five steps from around 600° to a maximum

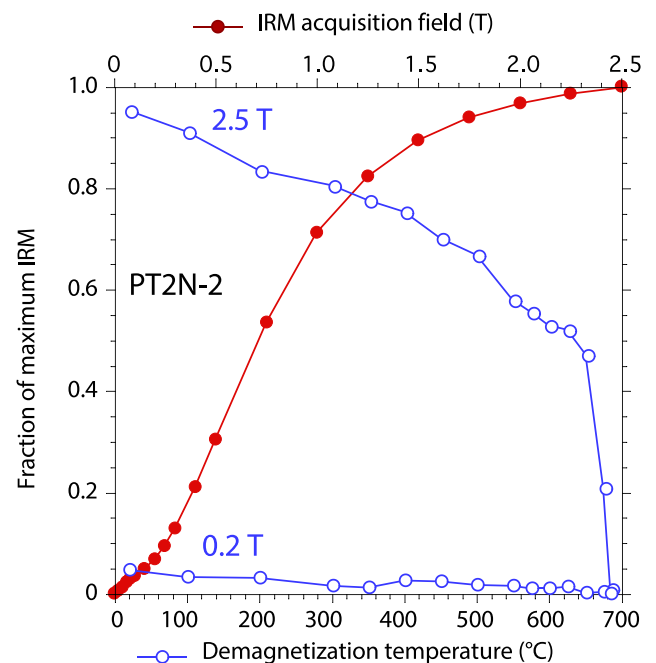


Fig. 3. Progressive acquisition in a sample from Sanaga sampling locality in the Ikakern Formation of an isothermal remanent magnetization (IRM) up to a direct field of 2.5 T along the specimen z-axis and subsequent thermal demagnetization of a composite orthogonal IRM after exposure to a direct field of 0.2 T along the specimen x-axis according to the method of Lowrie (1990). High coercivity of IRM and its unblocking concentrated between 660° and 680 °C point to hematite with a range of grain sizes as the predominant carrier of remanence.

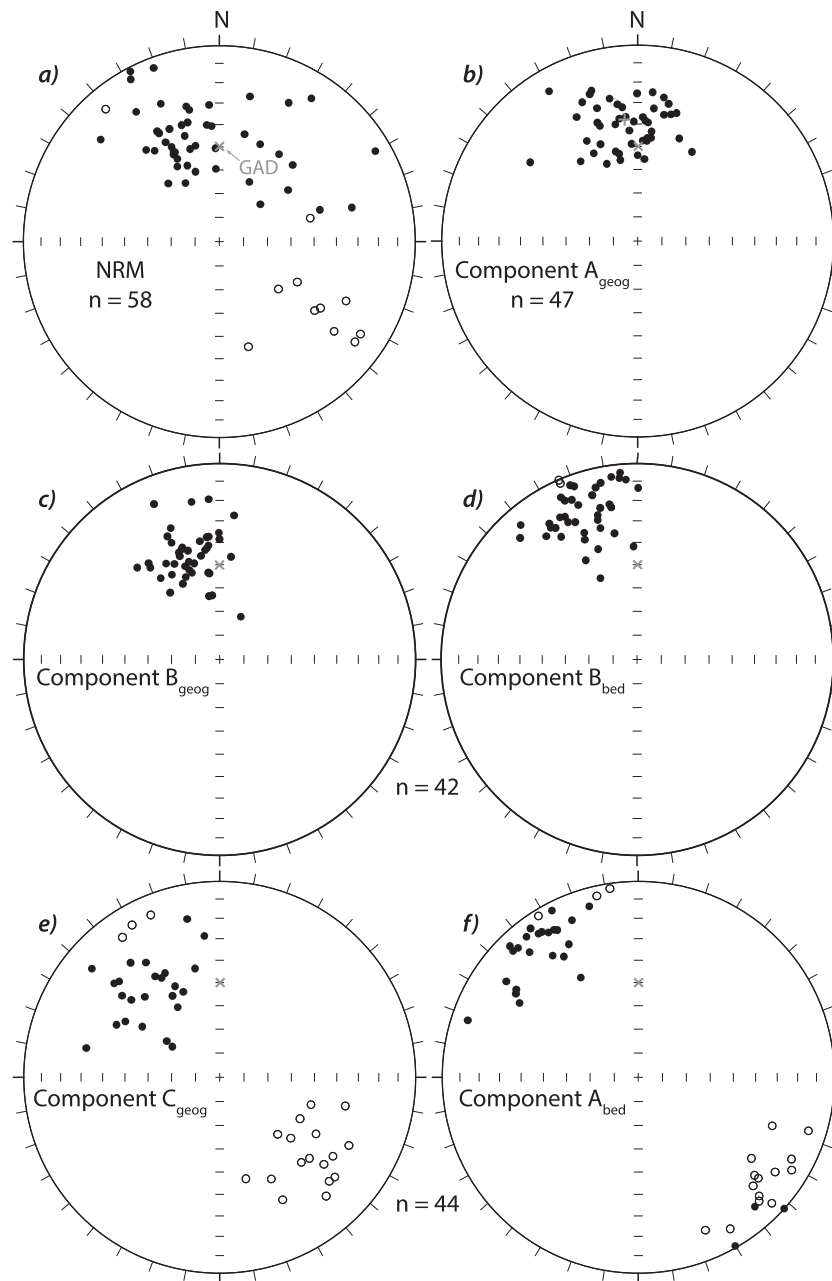


Fig. 4. Equal area plots (filled/open symbols on lower/upper hemisphere) of sample directions for a) NRM with geocentric axial dipole field direction at $D = 0^\circ$ $I = 50^\circ$ marked by a gray X in this and other plots; b) A component (100° – 300° C) with mean at $D = 353.9^\circ$ $I = 38.2^\circ$ ($a95 = 4.9^\circ$) marked by gray cross; c) B component (300° – 600° C) in geographic coordinates and d) in bedding coordinates; e) C component (600° – 675° C) in geographic coordinates and f) in bedding coordinates. See Tables 1 and 2 for mean directions by site of components B and C.

of 685° C anchored to the origin for component C) were calculated using standard least-squares methods (Kirschvink, 1980).

An attempt to extract a recently acquired viscous component A produced acceptable data (maximum angular deviation, MAD, of 15° or less) from 47 samples with a mean direction (Declination, $D = 353.9^\circ$, Inclination, $I = 38.2^\circ$, 95% confidence radius, $a95 = 4.9^\circ$) that is somewhat shallower than the geocentric axial dipole field direction for the sampling region ($D = 0^\circ$, $I \sim 50^\circ$) (Fig. 4B). Instead, the A component more closely resembles the intermediate unblocking temperature B component (Fig. 4C) that is also of uniform normal polarity and with which it may have an overlapping unblocking temperature spectrum. The mean B component for 42 samples from 11 sites with more than one accepted sample data in geographic coordinates is $D = 345.1^\circ$, $I = 42.7^\circ$ ($a95 = 6.4^\circ$) (Table 1), which is 8° but not significantly different from component A but more than 12° from the geocentric axial

dipole field. The mean B component direction becomes shallower with bedding corrections (Fig. 4D) but with an insignificant change in the precision parameter (Table 1).

The high unblocking temperature C component was isolated in 44 samples from the three sampling localities. The C component has a bipolar directional distribution with all but one sample from the sites at Ida Ou Kess Issar and Sanaga having northwesterly and down (normal polarity) directions (Fig. 5A, B, C) whereas samples from Si Mbark ou Mbarek have southeasterly and up (reverse polarity) directions (Fig. 5D) (Fig. 4E). Bedding attitudes at the Si Mbark ou Mbarek and Ida Ou Kess Issar localities are essentially homoclinal with dips typically around 15° – 35° down to the northwest so that the C component directions become shallower with bedding tilt correction (Fig. 4F). At the Sanaga locality, however, strata with opposing dips were sampled, which allows a local fold test to help constrain the origin of the

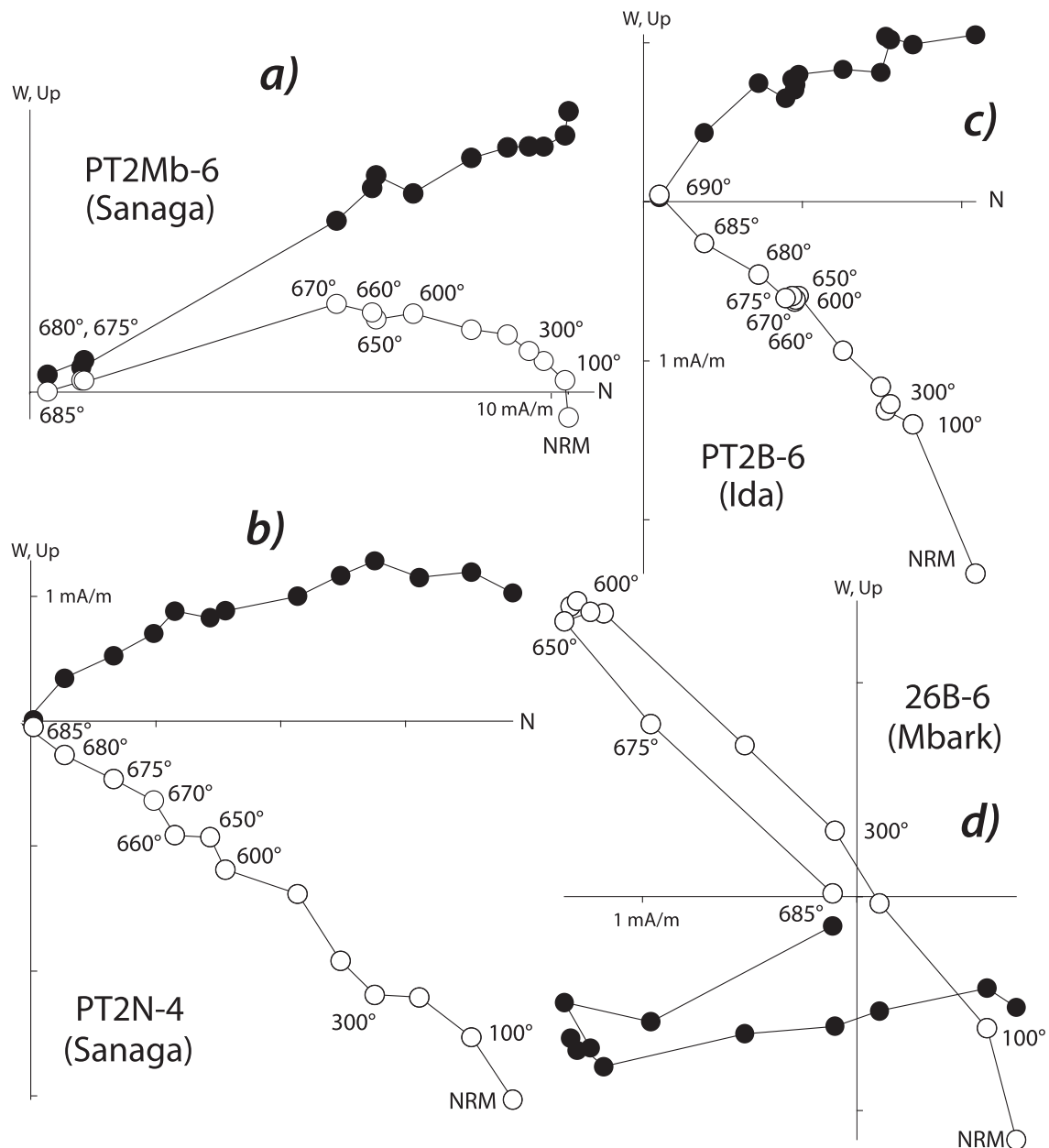


Fig. 5. Vector end-point plots of progressive thermal demagnetization of NRM of representative samples of Ikakern Formation from a) and b) Sanaga, c) Ida Ou Kess Issar, and d) Si Mbark ou Mbarek, all in geographic coordinates with open/closed symbols plotted on vertical/horizontal planes. Nominal endpoint temperatures for component A (100°, 300 °C), component B (300°, 600 °C) and all steps for component C (>600 °C) are labeled in each sample plot. Typical demagnetization schedule was 100°, 200°, 300°, 400°, 500°, 550°, 600°, 625 or 650°, 660°, 670°, 675° and sometimes 680°, 685° and 690 °C.

magnetizations. For the four sites at Sanaga (Table 2), the precision parameter, k , increases with corrections for bedding tilt from 6.4 to 50.2, a factor of 7.84 that far exceeds the 95% confidence threshold of 4.28 for a positive fold test (McElhinny, 1964). For all 14 sites, the overall precision parameter increases with corrections for bedding tilt by a factor of 1.91, just shy of 1.93 for 95% confidence.

The angular unconformity with the overlying Early to Late Triassic Timesgadiouine Formation constrains the folding and tilting of the Ikakern and hence acquisition of the C component with a positive local fold test to sometime between that and deposition. The reverse and normal polarity site mean characteristic magnetizations depart from being antipodal by 13.7°, which we interpret as due to residual contamination by a noncoaxial magnetization like the B component but whose contribution should get minimized by combining the site data of opposite polarities. The presence of reverse and normal

polarities combined with the positive local fold test point to acquisition in the latest part Kiaman Reverse Polarity Superchron, which ended at ~265 Ma (Lanci et al., 2013), and/or the ensuing Illawara mixed polarity interval in the late Permian, essentially 265–252 Ma. These age constraints build on the biostratigraphic context of the Ikakern Formation to indicate a nominal 260 Ma age for the C component.

In terms of calculating a more accurate paleomagnetic pole from the C component, there is strong and mounting evidence for the prevalence of inclination error (flattening) in early-acquired magnetizations carried by hematite in sedimentary red bed units such as the Ikakern (Bilardello and Kodama, 2010b; Kent and Olsen, 2008; Kent and Tauxe, 2005; Tauxe and Kent, 1984; Tauxe and Kent, 2004). Fortunately, there are techniques available to diagnose a systemic shallow bias, for example, using a measured magnetic anisotropy (Bilardello and Kodama, 2010a) or as in our case the distribution of the ChRM directions

Table 1
Overprint magnetization (B component) site mean directions for Ikakern Formation, Argana Basin, Morocco.

Site	sLat (°N)	sLon (°W)	n/N	k _g	Stk (°)	Dp (°)	gDec (°)	gInc (°)	bDec (°)	blnc (°)	nVGP _g	
											°Lon	°Lat
Si Mbark ou Mbarek												
^a 26A	30.8635	9.0133	1	–	213.5	33.0	5.7	27.1	356.4	9.0	152.1	72.7
26B	30.8640	9.0130	3	74.8	225.5	32.0	349.2	36.1	342.5	8.3	215.4	75.4
28A	30.8647	9.0242	2	27.8	212.5	14.0	346.3	52.5	336.6	41.6	275.4	78.2
28B	30.8667	9.0255	2	16.5	208.5	22.0	0.4	32.9	351.0	20.9	169.3	77.1
28C	30.8665	9.0277	2	300.7	219.5	27.0	340.2	49.6	328.5	22.9	264.6	73.0
Mean of site means:			4	49.1 (24.9)			349.84	43.0	339.9	23.6	234.1	70.2 (A95 = 12.3°)
Sanaga												
PT2Ma	30.7292	9.1040	4	107.4	295.5	27.0	341.0	41.6	351.4	25.6	243.6	71.8
PT2Mb	30.7293	9.1040	2	77.6	52.5	26.0	343.4	17.5	349.2	41.3	216.2	63.6
PT2N	30.7290	9.1040	5	85.0	243.5	35.0	342.6	43.1	340.1	7.6	245.0	73.6
PT2P	30.7295	9.1048	4	25.3	215.5	41.0	349.2	57.5	329.2	21.2	303.6	78.4
Mean of site means:			4	28.5 (20.9)			343.0	40.9	341.7	24.3	242.4	73.9 (A95 = 13.3°)
Ida Ou Kess Issar												
PT2A	30.6465	9.1877	6	62.3	232.5	27.0	335.4	43.2	332.3	12.7	252.7	67.6
PT2B	30.6503	9.1845	6	58.5	251.5	21.0	333.4	46.1	335.4	13.2	259.7	66.6
PT2L	30.6475	9.1878	5	39.3	244.5	33.0	353.7	46.3	348.1	3.6	233.6	83.7
Mean of site means:			3	102.6 (68.3)			340.7	45.6	338.7	19.0	253.5	72.8 (A95 = 15.4°)
Overall mean of site means:			11	43.9 (35.5)			345.1	42.7	340.4	22.2	242.8	75.8 (A95 = 6.0°)

Sampling sites of Ikakern Formation with GPS coordinates (sLat, sLon) and number of specimens (n, bold N for number of sites for locality and overall means) providing intermediate unblocking temperature (~300°–600 °C) B component directions whose site mean declinations and inclinations are given in geographic (gDec, gInc) coordinates and after correction for bedding attitude (strike, Stk, and dip, Dp, 90° clockwise from Stk) in bedding coordinates (bDec, blnc) with associated Fisherian precision parameter (k_g) for site means in geographic coordinates and also for locality and overall means based on sites where values in parentheses are for bedding coordinates; none of the precision parameters for the locality and overall means changes significantly after bedding corrections. North virtual geomagnetic poles (nVGP_g) calculated from site mean ChRM directions in geographic coordinates (gDec, gInc); locality mean and overall mean nVGP_g calculated from site nVGP_g.

^a Site with only one accepted sample data is not used for locality or overall mean calculations.

Table 2
Characteristic magnetization (C component) site mean directions for Ikakern Formation (Argana Basin, Morocco).

Site	sLat (°N)	sLon (°W)	n/N (°)	k _b (°)	Stk (°)	Dp (°)	gDec (°)	gInc (°)	bDec	blnc	nVGP _{f=0.6}	
											°Lon	°Lat
Si Mbark ou Mbarek												
26A	30.8635	9.0133	2	88.5	213.5	33.0	128.5	–41.4	127.3	–8.4	246.5	35.4
26B	30.8640	9.0130	2	20.8	225.5	32.0	144.7	–30.9	143.4	0.8	225.7	43.1
26C	30.8623	9.0117	3	21.4	223.5	33.0	152.2	–44.6	145.8	–12.3	235.5	52.2
26D	30.8623	9.0117	3	27.5	234.5	24.0	137.7	–38.1	130.9	–14.9	250.1	41.3
26E	30.8655	9.0207	2	200.0	255.5	13.0	131.8	–27.7	134.6	–16.7	249.4	45.2
28A	30.8647	9.0242	2	237.6	212.5	14.0	121.3	–28.4	121.4	–14.4	255.4	33.0
^a 28B	30.8667	9.0255	1	–	208.5	22.0	106.6	–49.8	109.9	–28.2	273.5	28.4
28C	30.8665	9.0277	3	30.5	219.5	27.0	117.2	–50.2	119.6	–21.2	262.6	34.1
Mean of site means:			7	44.9 (39.6)			132.4	–37.9	132.0	–12.6	247.2	41.2 (A95 = 8.7°)
Sanaga												
PT2Ma	30.7292	9.1040	4	29.0	295.5	27.0	309.0	35.6	324.2	25.6	253.9	56.8
PT2Mb	30.7293	9.1040	3	183.4	52.5	26.0	330.8	–12.3	330.8	13.5	231.6	56.7
PT2N	30.7290	9.1040	5	84.0	243.5	35.0	328.0	42.5	329.4	7.6	226.8	52.3
PT2P	30.7295	9.1048	2	49.6	215.5	41.0	318.3	61.8	311.9	21.2	256.0	44.7
Mean of site means:			4	50.2 ^b (6.4)			322.3	33.0	324.2	17.1	242.6	53.4 (A95 = 12.3°)
Ida Ou Kess Issar												
PT2A	30.6465	9.1877	6	12.9	232.5	27.0	321.3	39.7	321.6	12.7	239.9	49.2
PT2B	30.6503	9.1845	4	58.6	251.5	21.0	311.7	31.8	315.7	13.2	245.1	44.6
PT2L	30.6475	9.1878	2	56.9	244.5	33.0	350.8	35.5	347.7	3.6	196.0	60.0
Mean of site means:			3	21.2 (22.8)			327.7	36.8	328.4	10.1	230.6	53.1 (A95 = 27.2°)
Overall mean of site means:			14	32.3 (16.9)			318.2	36.5	319.0	13.5	246.3	48.8 (A95 = 7.2°)

Sampling sites of Ikakern Formation at three localities with GPS coordinates (sLat, sLon) and number of samples (n, bold N for number of sites for locality and overall means) providing high unblocking temperature (~600°–675 °C) C component (ChRM) directions whose site mean declinations and inclinations are given in geographic coordinates (gDec, gInc) and after correction for bedding attitude (strike, Stk, and dip, Dp, 90° clockwise from Stk) in bedding coordinates (bDec, blnc) with associated Fisherian precision parameter (k_b) for site means and also for locality and overall site means where values in parentheses are for geographic coordinates. North virtual geomagnetic poles (nVGP_{f=0.6}) calculated from site mean ChRM directions in bedding coordinates (bDec, blnc) adjusted by an inclination flattening factor, f = 0.6; locality mean and overall mean nVGP_{f=0.6} calculated from site nVGP_{f=0.6}.

^a Site with only one accepted sample data is not used for locality or overall mean calculations.

^b Precision parameter for Sanaga in bedding coordinates is a factor of 7.84 greater than in geographic coordinates that well exceeds test value of 4.28 for n = 4 at 95% confidence (McElhinny, 1964), indicating a significant (positive) local fold test.

using the Elongation/Inclination (EI) technique (Tauxe and Kent, 2004). Although the number of samples available for analysis is less than desirable, application of the EI procedure on the 44 tilt-corrected ChRM sample directions (Fig. 6) derives a flattening factor, $f = 0.48$, consistent with either a depositional detrital remanence or an early acquired chemical remanence subject to compaction effects before the sediments became thoroughly lithified. The derived value, however, is somewhat low compared to the range of flattening factors determined in laboratory redeposited and ancient red bed units analyzed in the references cited above and may be influenced by deviations from antipodality mainly in declination of the normal and reverse polarity directions. Accordingly, we adopt a nominal f -value of 0.6 that has been summarily used in some compilations of published poles from clastic sedimentary rocks (e.g., Torsvik et al., 2012) to calculate a mean pole for the Ikakern based on corrected site mean virtual geomagnetic poles; the mean pole is located at $47.4^\circ\text{N } 242.9^\circ\text{E } A95 = 6.9^\circ$ (Table 2). For comparison, the mean pole corrected using $f = 0.48$ from the EI analysis would be at

$48.8^\circ\text{N } 246.3^\circ\text{E } A95 = 7.2^\circ$ whereas at $44.7^\circ\text{N } 237.8^\circ\text{E } A95 = 6.6^\circ$ for $f = 1$ or no flattening (Fig. 7a).

The dual polarity pre-folding Ikakern pole corrected for inclination flattening strengthens the mean 260 Ma pole for Gondwana, which was based on six results deemed reliable spanning 252–267 Ma in a centered 20 Myr time window (from NW Africa, Adria, Southern Africa, Australia, and two from South America; Kent and Muttoni, 2020); the addition of the Ikakern pole changes it by less than 1° while narrowing the uncertainty ($52.2^\circ\text{N } 239.8^\circ\text{E } A95 = 5.0^\circ$ $N = 7$, in NW Africa coordinates; Table 3).

4. Implications for configuration of Pangea in the Late Permian

The Ikakern B and C component poles are compared in Fig. 7b to the Gondwana APW reference path in NW Africa coordinates (Table 4). The B component pole ($75.8^\circ\text{N } 242.8^\circ\text{E } A95 = 8.0^\circ$) is only 4° away from the 200 Ma mean Gondwana pole and more specifically,

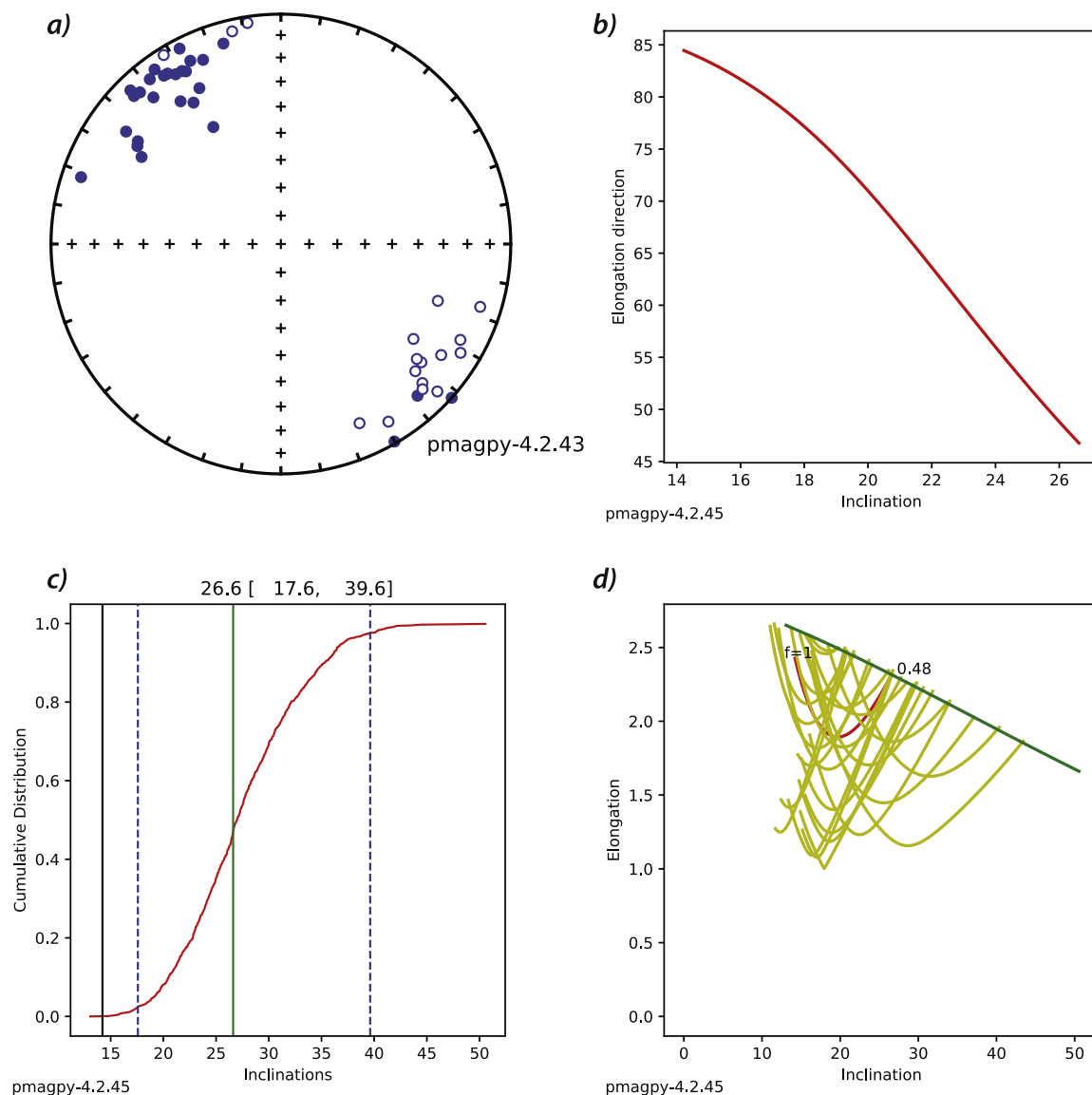


Fig. 6. EI analysis (Tauxe and Kent, 2004) of C component (ChRM) directions of Ikakern Formation. a) ChRM sample directions corrected with $f = 0.48$ in bedding coordinates; b) modeled elongation versus inclination for TK03.GAD field model; c) cumulative distribution of ChRM inclinations corrected with $f = 0.48$ in bedding coordinates with mean and 95% confidence limits shown by solid green line and dashed lines respectively; d) trajectories of sample ChRM inclinations in bedding coordinates by varying f from 1 to intersection of expected EI relationship from TK03.GAD field model; best-fit for this population was estimated as $f = 0.48$. Analysis and plots generated by PmagPy (Tauxe et al., 2016). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

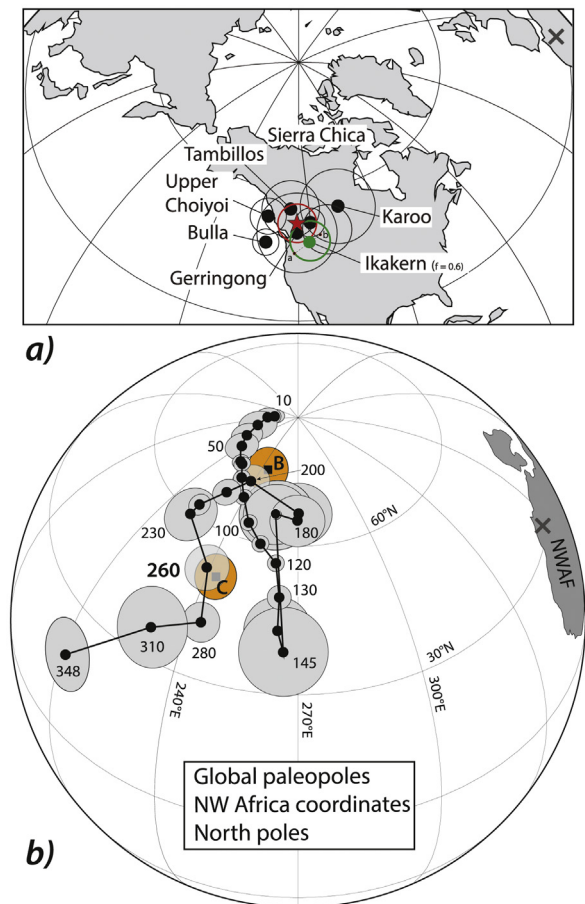


Fig. 7. a) Paleomagnetic (north) poles deemed reliable for Late Permian rock units from various areas in Gondwana in NW Africa coordinates (from Kent and Muttoni, 2020). Sampling locality of Ikakern Formation in Morocco indicated by X. Pole for Ikakern C component with flattening factor, $f = 0.6$, is plotted as green circle and is bracketed by pole with no correction ($f = 1$) labeled 'a' that is joined by dashed line segment of great circle to sampling site to pole with correction for $f = 0.48$ labeled 'b'. Overall 260 Ma mean pole for Gondwana is indicated by red star (with 95% confidence circle also in red) and includes Ikakern pole ($f = 0.6$) (Table 3). b) APW path for Gondwana in NW Africa coordinates from 348 Ma (Early Carboniferous) to 10 Ma (Miocene) with mean north poles generally centered on 20 Myr windows (Table 4). Sampling locality of Ikakern Formation in Morocco indicated by X. Pole positions marked by filled squares with A95 circles in orange are for Ikakern B component in geographic coordinates (labeled 'B') and C component in bedding coordinates (labeled 'C'). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

6° from the mean pole associated with igneous units of CAMP activity at ~200 Ma in NW Africa, which for six studies give a mean pole at 69.7°N 240.4°E $A95 = 7.9^\circ$ (Kent and Irving, 2010). This correspondence in directions and common normal polarity point to thermochemical overprinting associated with local CAMP activity as a plausible agent of the B component overprint. As mentioned above in the geologic setting, outcrops of CAMP igneous intrusions occur near the Sanaga and Ida Ou Kess Issar localities; moreover, the several thousand meters of Argana basin section between the Ikakern and CAMP lava flows (Fig. 2) would have produced burial temperatures further conducive to alteration. We regard the agreement of the B component with the CAMP mean pole as showing that this area of Morocco was tectonically coherent with NW Africa since at least ~200 Ma.

Alternatively, the Ikakern may have been affected by a presumed overprint described in Jurassic marine carbonates from the High Atlas Mountains bordering the Argana basin to the southeast and attributed to a tectonic event at 100 Ma on the basis of the magnetization direction (pole position at ~62°N 244°E calculated from a mean direction in 100

sites of $D = 332.2^\circ I = 34.5^\circ$ given in Calvin et al. (2018)). The B component pole is a significant 14° away from the High Atlas overprint pole. Moreover, it is interesting to note that the APW reference path loops around (Fig. 7b) such that magnetizations corresponding in direction to the 100 Ma (mid-Cretaceous) mean pole may not be readily distinguishable from those corresponding to the 160 Ma and 170 Ma (mid-Jurassic) mean poles that are closer to the depositional age of the Jurassic carbonates.

The 260 Ma mean poles for Gondwana (now reinforced by the Ikakern result) and Laurasia that are based on igneous or inclination-flattening corrected sedimentary results (Kent and Muttoni, 2020) are compatible with a standard Pangea A-like configuration in the Late Permian, with the northern margin of Africa against the eastern margin of North America (Fig. 8a). Paleomagnetic data from Gondwana and Laurasia for earlier time periods (Carboniferous–Early Permian), however, produce a fit that is much too tight for Pangea A and instead favor a configuration known as Pangea B (Morel and Irving, 1981) with the northern margin of South America instead against the eastern margin of North America. Although the tectonic transition to a Pangea A-type configuration was initially presumed to have occurred mainly in the Triassic and completed just prior to Pangea breakup in the early Jurassic (Irving, 1977; Irving, 1981; Morel and Irving, 1981), subsequent analyses of paleomagnetic and tectonic data that entertained Pangea B (e.g., Muttoni et al., 1996, 2003, 2009; Şengör, 2016; Kent and Muttoni, 2020) place the transition within the Permian, a timing eventually accepted by the main proponent (Irving, 2004). A Pangea A-type configuration in the Late Permian has thus become almost universally accepted. Nonetheless, a recently published paleomagnetic study of the upper El Centinelo volcanic complex in Argentina advocated the persistence of Pangea B through the Late Permian (Tomezzoli et al., 2018) (Fig. 8b).

The global APW path with respect to South America coordinates using the same mean poles as for NW Africa but as south poles for convenience of this discussion is shown in Fig. 8c and listed in Table 4. South America contributes two of the six (now seven) reference poles deemed reliable to estimate the Late Permian (260 Ma) paleopole position for Gondwana (Kent and Muttoni, 2020). These South America poles are from the Sierra Chica volcanics with U–Pb zircon dates of $263.0 \pm 1.6 / -2.0$ Ma (Domeier et al., 2011a) and the Upper Choyoi Group with U–Pb and ^{40}Ar – ^{39}Ar dates giving an estimated age at ~264 Ma and supported by a positive fold test (Domeier et al., 2011b); both units are in the Gondwanide mobile belt of Argentina (Kay et al., 1989). In contrast, the El Centinelo volcanics, which are also in the Gondwanides of Argentina, produced more disparate results (Tomezzoli et al., 2018). The El Centinela-II pole (69.2°S 48.2°E $A95 = 3.5^\circ$) from the upper part of a ~150 m-thick section with a 276 ± 11 Ma U–Pb zircon date is widely different from both the 260 Ma (by $20.0 \pm 6.1^\circ$) as well as the 280 Ma (by $14.6 \pm 5.3^\circ$) mean Gondwana poles (Fig. 8c; Table 4). On the other hand, the El Centinela-I pole (60.8°S 356.6°E $A95 = 4.5^\circ$) from the lower part of the El Centinela section and indirectly assigned an age of 295.5 ± 8.0 Ma (Tomezzoli et al., 2018) falls more compatibly in between the 280 Ma (Early Permian) and 310 Ma (Late Carboniferous) mean Gondwana south poles in South American coordinates (Fig. 8c; Table 4), which happen to support Pangea B (Kent and Muttoni, 2020).

The El Centinela-II volcanics pole is thus anomalous and may reflect a remagnetization (perhaps in the Late Cretaceous–Early Jurassic based on where pole position falls on the APW path?), which some reports suggest may be repeated and widespread in South America (Bilardello et al., 2018; Font et al., 2012), and/or local tectonics such as rotations of 30° or more reported for Patagonia in the Late Permian (Luppo et al., 2019). There are many other poles reported from latest Paleozoic and Mesozoic rocks in South America discussed in (Tomezzoli et al., 2018) and references cited therein and above; here we would just suggest that differentiating ages of magnetization based on the Gondwana APW path in South America coordinates has an inherent difficulty

Table 3

Late Permian paleomagnetic reference poles from igneous rock units and inclination flattening-corrected sedimentary units for Laurasia (mean only given for 260 Ma) and for Gondwana.

#	Plate	Age (Ma)	Unit	Lat (°N)	Lon (°E)	A95 (°)	Reference
<i>Laurasia</i>							
	LAUR	260	Mean Late Permian (N = 10)	54.7	147.7	3.3	Kent and Muttoni (2020)
<i>Gondwana</i>							
67	ADR	252	Upper Bellerophon&Werfen	47.5	228.9	3.0	Scholger et al. (2000); Muttoni et al. (2003)
68	SAM	263	Sierra Chica, La Pampa	52.4	244.3	3.3	Domeier et al. (2011a)
69	SAM	264	Upper Choiyoi, Argentina	53.4	228.4	4.1	Domeier et al. (2011b)
70	AUS	265	Gerringong volcanics	50.1	239.4	9.1	Belica et al. (2017)
71	SAF	266.5	Karoo redbeds, S. Africa	55.3	255.7	8.5	Lanci et al. (2013)
72	SAM	267	Tambillos Fm., Argentina	55.5	236.8	6.5	Rapalini and Vilas (1991)
	NWAF	260	Ikakern Formation, Morocco	48.8	246.3	7.2	This Study
	GOND	260	Mean (N = 6: #67–72)	52.7	238.6	5.7	Kent and Muttoni (2020)
	GOND	260	Mean (N = 7: #67–72 + Ikakern)	52.2	239.8	5.0	This Study

= paleopole item number (Kent and Muttoni, 2020). **Plate:** Laurasia (LAUR) is comprised of data from North America and Stable Europe. Gondwana (GOND) is comprised of data from ADR, Adria (=NW Africa); AUS, Australia; NEAF, NE Africa; NWAF, NW Africa; SAM, South America. **Age:** mean age of unit. **Unit:** sampled lithological units all comprised of igneous rocks except for entry #67, Bulla section sediments of Scholger et al. (2000) EI-corrected for inclination shallowing by Muttoni et al. (2013); entry #71, sediments EI-corrected for inclination shallowing by Lanci et al. (2013); and the EI corrected Ikakern Formation of this study. **Lat** and **Lon** are latitude (°N) and longitude (°E) of paleopoles in European coordinates for Laurasia and of paleopoles rotated to NW Africa for Gondwana using rotation parameters of Lottes and Rowley (1990). **A95:** radius cone of 95% confidence. **Reference:** primary reference to paleopole; see also Kent and Muttoni (2020).

Table 4

Preferred APW path for Gondwana for the Late Paleozoic, Mesozoic and Cenozoic in NW Africa and South America coordinates.

Age (Ma)	A95 (°)	NW Africa		S. America		Ref1	Ref2
		Lat (°N)	Lon (°E)	Lat (°S)	Lon (°E)		
10	2.0	85.3	173.5	85.9	331.0	K&M2013	B&C2003
20	2.7	83.9	175.9	84.7	313.8	K&M2013	B&C2003
30	3.9	81.8	190.7	83.7	312.6	K&M2013	B&C2003
40	3.2	79.0	201.1	82.6	319.2	K&M2013	B&C2003
50	3.4	76.9	210.3	82.1	321.8	K&M2013	B&C2003
60	1.7	74.4	221.9	82.4	342.8	K&M2013	B&C2003
70	1.7	74.2	224.4	83.2	337.7	K&M2013	K&I2010
80	1.7	71.7	231.3	83.3	343.5	K&M2013	K&I2010
90	1.7	68.2	239.5	84.8	341.1	K&M2013	K&I2010
100	1.7	63.5	247.4	86.0	358.1	K&M2013	K&I2010
110	1.7	59.8	254.9	87.9	37.6	K&M2013	K&I2010
120	1.7	56.3	262.0	86.9	104.3	K&M2013	K&I2010
130	2.4	49.5	264.2	79.7	61.9	M + 2013	K&I2010
140	6.8	42.8	264.4	72.0	55.0	M + 2013	K&I2010
145	9.0	38.6	266.2	67.6	55.8	M + 2013	K&I2010
160	7.5	66.4	259.1	83.4	202.8	M + 2013	K&I2010
170	6.5	66.5	258.7	83.3	204.3	M + 2013	K&I2010
180	5.5	65.4	269.7	81.7	170.6	M + 2013	K&I2010
190	6.7	66.9	270.3	80.5	177.7	M + 2013	K&I2010
200	3.8	71.9	238.1	76.2	237.5	M + 2013	K&I2010
210	2.9	67.5	229.5	76.3	259.7	M + 2013	K&I2010
220	2.3	62.3	222.4	73.8	281.4	M + 2013	K&I2010
230	5.7	59.7	222.0	73.2	290.2	M + 2013	K&I2010
260	5.0	52.2	239.8	78.2	338.2	K&M2020 + this study	
280	4.0	41.4	243.2	69.9	5.6	K&M2020	
310	8.2	36.2	230.6	61.2	348.5	K&M2020	
348	7.5	18.8	211.2	38.0	336.5	K&M2020	

Age is center of 20 Myr averaging window stepped every 10 Myr (except between 140 and 160 Ma, 230–260 Ma, and 280–310 Ma; pole at 348 Ma is based on only one entry). Mean (north) poles for Gondwana in NW Africa coordinates and same data as south poles in South America coordinates are listed as latitude (Lat), longitude (Lon), and radius of 95% circle of confidence (A95). Ref1 are references for mean poles and rotation parameters (K&M2013 = (Kent and Muttoni, 2013); M + 2013 = Muttoni et al., 2013; K&M2020 = Kent and Muttoni, 2020). Ref2 are references for additional pole listings (B&C2003 = Besse and Courtillot, 2002(erratum); K&I2010 = Kent and Irving, 2010, with rotation parameters in Table 5.

because the APW path has rather limited range and the mean poles are within 15° of the geocentric axial dipole axis (for 260–160 Ma mean poles, the mean latitude is 78.5° ± 3.9° S.D.) (Fig. 8c). In contrast, the Gondwana APW path in NW Africa coordinates has a broader latitudinal

range typically more than 25° from the geocentric axial dipole axis (for 260–160 Ma mean poles, the mean latitude is 64.3° ± 5.6° S.D.) (Fig. 7b) (Table 4), making it easier to discriminate paleomagnetic directions from more recent remagnetizations although the APW path does loop back on itself as discussed for the putative 100 Ma remagnetization reported in the carbonates from the South Atlas Mountains. In any case, the anomalous location of the El Centinela-II pole with respect to the Gondwana APW path projected to South America suggests that this and nearby poles from a complicated tectonic setting may not be appropriate for constraining continental reconstructions.

Table 5

Rotation parameters to transfer mean paleopoles for 130 to 230 Ma (Kent and Irving, 2010) from NAM coordinates to SAM coordinates.

Magnetic Anomaly	Age Ma	Euler pole			Reference
		Lat °N	Lon °E	Angle	
NAM to SAM rotation poles					
M11	131.0 Ma	66.1	341.3	58.0	Roest et al. (1992)
M16	139.2	66.2	341.7	59.7	Roest et al. (1992)
M21	148.5	66.2	341.7	62.1	Roest et al. (1992)
M25	154.2	66.7	344.1	64.9	Roest et al. (1992)
BSMA	170	67.0	346.8	72.1	Klitgord and Schouten (1986)
Closure	175	67.0	348.0	75.6	Klitgord and Schouten (1986)
Fit	180–252	66.4	345.9	75.1	Lottes and Rowley (1990)
Interpolated NAM to NWAF rotation poles:					
130 Ma		66.1	341.3	58.0	
140		66.2	341.7	60.0	
145		66.2	341.7	61.3	
160		66.8	345.2	67.6	
170		67.0	346.8	72.1	
180		66.4	345.9	75.1	
190		66.4	345.9	75.1	
200		66.4	345.9	75.1	
210		66.4	345.9	75.1	
220		66.4	345.9	75.1	
230		66.4	345.9	75.1	
Interpolated NAM to SAM rotation poles:					
130 Ma		38.56	75.62	15.67	
140		42.64	71.05	16.99	
145		45.13	68.09	17.83	
160		50.19	56.62	23.45	
170		52.55	50.02	27.51	
180–230		54.68	43.22	29.66	
NWAF to SAM rotation pole:					
130–348 Ma		53.0	325.0	–51.0	Lottes and Rowley (1990)

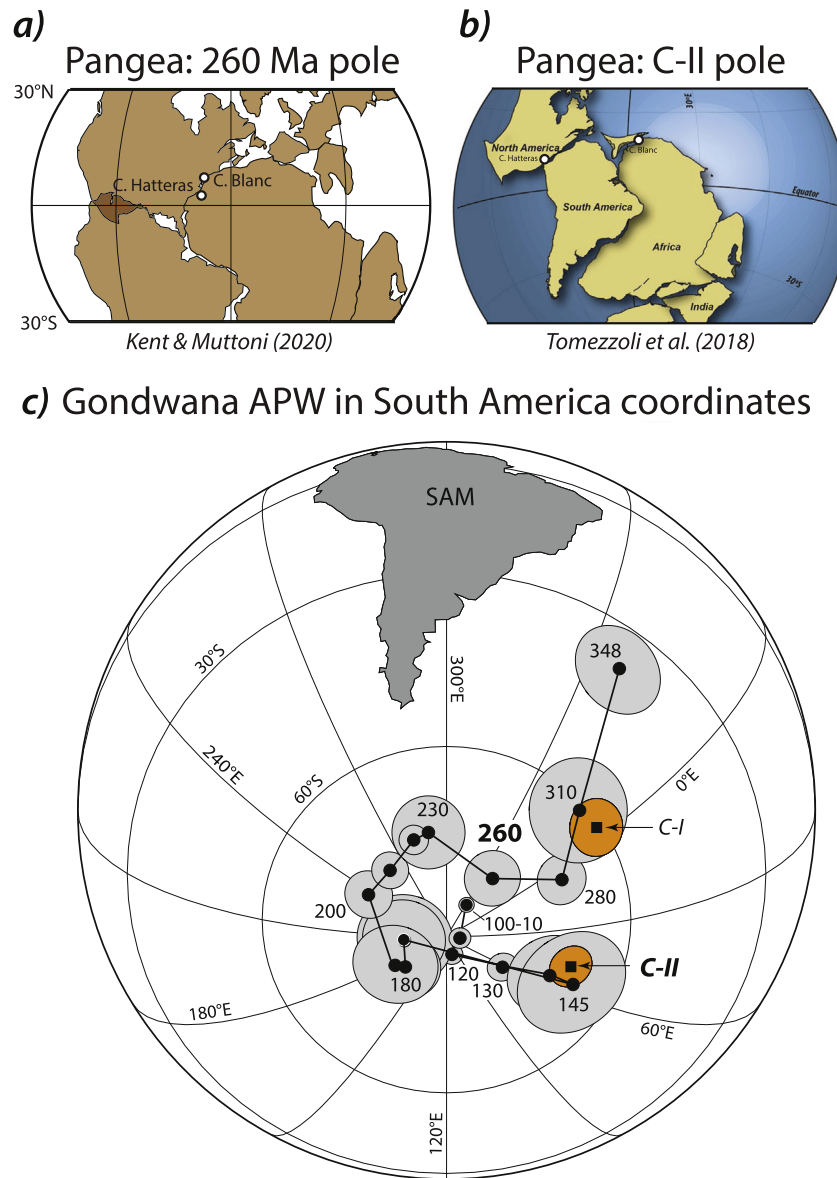


Fig. 8. a) APW path for Gondwana in South America coordinates from 348 Ma (Early Carboniferous) to 10 Ma (Miocene) with mean poles generally centered on 20 Myr windows (Table 4). C-I and C-II are El Centinela poles from Gondwanides of Argentina (Tomezzoli et al., 2018). b) Pangea A for Late Permian according to mean 260 Ma mean pole for Gondwana (and Laurasia) (Kent and Muttoni, 2020). c) Pangea B for Late Permian according to C-II pole (Tomezzoli et al., 2018).

5. Conclusions

Paleomagnetic analysis of samples from the Ikakern Formation provide a new reference pole for the Late Permian of NW Africa that is supported by a positive local fold test and a constraint on inclination flattening using the EI technique. The Ikakern pole is concordant with the most reliable igneous and inclination-flattening corrected sedimentary poles of nominal 260 Ma age from the independently reconstructed Gondwana continents and in conjunction with reliable igneous and inclination-flattening corrected sedimentary poles of comparable nominal 260 Ma age from Laurasia, are compatible with a classic Pangea A configuration in the Late Permian. Some recently published poles from South America that were used to argue for a Pangea B configuration persisting through the Late Permian are inconsistent with the Gondwana mean pole and may reflect remagnetizations or tectonic rotations in their source areas subject to further studies. Transferring reliable poles from other Gondwana continents using independent

reconstruction parameters may be a useful tool to evaluate the timing, severity and extent of localized magnetic resetting and changes in structural reference frame.

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Corrigendum

Corrigendum to “A Late Permian paleopole from the Ikakern Formation (Argana basin, Morocco) and the configuration of Pangea” [Gondwana Res., 92, 2021, 266–278]



Dennis V. Kent^{a,b,*}, Paul E. Olsen^a, Giovanni Muttoni^c, Mohammed Et-Touhami^d

^a Lamont-Doherty Earth Observatory of Columbia University, Palisades, NY 10964, USA

^b Earth and Planetary Sciences, Rutgers University, Piscataway, NJ 08854, USA

^c Dipartimento di Scienze della Terra 'Ardito Desio', Università degli Studi di Milano, via Mangiagalli 34, I-20133 Milan, Italy.

^d 2GPMH, Département des Sciences de la Terre, Université Mohammed Premier Oujda, 60000 Oujda, Morocco

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The authors would like to correct in the body of **Table 5** a section heading, which should read NAM to NWAF (rather than SAM) rotation poles, and the placement of the section NWAF to SAM rotation pole, which should come just above Interpolated NAM to SAM rotation poles. The corrected **Table 5** is given below. And of minor significance to this paper but reported here for consistency, a correction for an error in the entry of longitude for the 145 Ma Swartruggen-Bumbeni pole in Table 3 of Kent and Irving (2010), as noted in [Fu et al. \(2020\)](#), results in a revised 145 Ma mean pole in **Table 4** with A95 = 15.2° and located at 38.8°N, 269.8°E in NW Africa coordinates and at 67.1°S, 63.0°E in S. America coordinates. These pole positions differ by 2.8° from those originally tabulated for 145 Ma in **Table 4**.

The authors regret these errors.

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Table 5

Rotation parameters to transfer mean paleopoles for 130 to 230 Ma (Kent and Irving, 2010) from NAM coordinates to SAM coordinates.

Magnetic Anomaly	Age (Ma)	Euler pole			Reference
		Lat (°N)	Lon (°E)	Angle (°cw)	
NAM to NWAF rotation poles					
M11	131.0	66.1	341.3	58.0	Roest et al. [1992]
M16	139.2	66.2	341.7	59.7	Roest et al. [1992]
M21	148.5	66.2	341.7	62.1	Roest et al. [1992]
M25	154.2	66.7	344.1	64.9	Roest et al. [1992]
BSMA	170.0	67.0	346.8	72.1	Klitgord and Schouten [1986]
Closure	175.0	67.0	348.0	75.6	Klitgord and Schouten [1986]
Fit	180–252	66.4	345.9	75.1	Lottes and Rowley [1990]
Interpolated NAM to NWAF rotation poles					
	130	66.1	341.3	58.0	
	140	66.2	341.7	60.0	
	145	66.2	341.7	61.3	
	160	66.8	345.2	67.6	
	170	67.0	346.8	72.1	
	180	66.4	345.9	75.1	
	190	66.4	345.9	75.1	
	200	66.4	345.9	75.1	
	210	66.4	345.9	75.1	
	220	66.4	345.9	75.1	
	230	66.4	345.9	75.1	
NWAF to SAM rotation pole					
	130–348	53.0	325.0	–51.0	Lottes and Rowley [1990]
Interpolated NAM to SAM rotation poles					
	130	38.6	75.6	15.7	
	140	42.6	71.1	17.0	
	145	45.1	68.1	17.8	
	160	50.2	56.6	23.5	
	170	52.6	50.0	27.5	
	180–230	54.7	43.2	29.7	

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* Corresponding author.

E-mail address: dvk@ldeo.columbia.edu (D.V. Kent).

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