

# Dietary options and behavior suggested by plant biomarker evidence in an early human habitat

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Edited by Thure E. Cerling, University of Utah, Salt Lake City, UT, and approved January 21, 2016 (received for review April 13, 2015)

**The availability of plants and freshwater shapes the diets and social behavior of chimpanzees, our closest living relative. However, limited evidence about the spatial relationships shared between ancestral human (hominin) remains, edible resources, refuge, and freshwater leaves the influence of local resources on our species' evolution open to debate. Exceptionally well-preserved organic geochemical fossils—biomarkers—preserved in a soil horizon resolve different plant communities at meter scales across a contiguous 25,000 m<sup>2</sup> archaeological land surface at Olduvai Gorge from about 2 Ma. Biomarkers reveal hominins had access to aquatic plants and protective woods in a patchwork landscape, which included a spring-fed wetland near a woodland that both were surrounded by open grassland. Numerous cut-marked animal bones are located within the wooded area, and within meters of wetland vegetation delineated by biomarkers for ferns and sedges. Taken together, plant biomarkers, clustered bone debris, and hominin remains define a clear spatial pattern that places animal butchery amid the refuge of an isolated forest patch and near freshwater with diverse edible resources.**

biomarker | leaf wax | carbon isotope | paleoecology | human evolution

**S**patial patterns in archaeological remains provide a glimpse into the lives of our ancestors (1–5). Although many early hominin environments are interpreted as grassy or open woodlands (6–8), fossil bones and plant remains are rarely preserved together in the same settings. As a result, associated landscape reconstructions commonly lack coexisting fossil evidence for hominins and local-scale habitat (microhabitat) that defined the distribution of plant foods, refuge, and water (7). This problem is exacerbated by the discontinuous nature and low time resolution often available across ancient soil (paleosol) horizons, including hominin archaeological localities. One notable exception is well-time-correlated 1.8-million-y-old paleosol horizons exposed at Olduvai Gorge. Associated horizons contain exceptionally preserved plant biomarkers along with many artifacts and fossilized bones. Plant biomarkers, which previously revealed temporal patterns in vegetation and water (8), are well preserved in the paleosol horizon and document plant-type spatial distributions that provide an ecosystem context (9, 10) for resources that likely affected the diets and behavior of hominin inhabitants.

Plant biomarkers are delivered by litter to soils and can distinguish plant functional type differences in standing biomass over scales of 1–1,000 m<sup>2</sup> (11). Trees, grasses, and other terrestrial plants produce leaf waxes that include long-chain *n*-alkanes such as hentriacontane (*n*C<sub>31</sub>), whereas aquatic plants and phytoplankton produce midchain homologs (e.g., *n*C<sub>23</sub>) (12, 13). The ratio of shorter- versus long-chain *n*-alkane abundances distinguish relative organic matter inputs from aquatic versus terrestrial plants to sediments (13):

$$P_{aq} = (nC_{23} + nC_{25}) / (nC_{23} + nC_{25} + nC_{29} + nC_{31})$$

Sedges and ferns are prolific in many tropical ecosystems (14). These plants both have variable and therefore nondiagnostic *n*-alkane profiles. However, sedges produce distinctive phenolic

compounds [e.g., 5-*n*-tricosylresorcinol (*n*R<sub>23</sub>)] and ferns produce distinctive midchain diols [e.g., 1,13-dotriacontanediol (C<sub>32</sub>-diol)] (*SI Discussion*).

Lignin monomers provide evidence for woody and nonwoody plants. This refractory biopolymer occurs in both leaves and wood, serves as a structural tissue, and accounts for up to half of the total organic carbon in modern vegetation (11). Lignin is composed of three phenolic monomer types that show distinctive distributions in woody and herbaceous plant tissues. Woody tissues from dicotyledonous trees and shrubs contain syringyl (*S*) and vanillyl (*V*) phenols (12), whereas cinnamyl (*C*) phenols are exclusively found in herbaceous tissues (12). The relative abundance of *C* versus *V* phenols (*C/V*) is widely used to distinguish between woody and herbaceous inputs to sedimentary and soil organic matter (15).

Plant biomarker <sup>13</sup>C/<sup>12</sup>C ratios (expressed as δ<sup>13</sup>C values) are sensitive indicators of community composition, ecosystem structure, and climate conditions (8). Most woody plants and forbs in eastern Africa use C<sub>3</sub> photosynthesis (6), whereas arid-adapted grasses use C<sub>4</sub> photosynthesis (8, 14). These two pathways discriminate differently against <sup>13</sup>C during photosynthesis, resulting in characteristic δ<sup>13</sup>C values for leaf waxes derived from C<sub>3</sub> (about −36.0‰) and C<sub>4</sub> (−21.0‰) plants (16). Carbon isotopic abundances of phenolic monomers of lignin amplify the C<sub>3</sub>–C<sub>4</sub> difference and range between *ca.* −34.0‰ (C<sub>3</sub>) and −14.0‰ (C<sub>4</sub>) in tropical ecosystems (15). Terrestrial C<sub>3</sub> plant δ<sup>13</sup>C values decrease with increased exposure to water, respired CO<sub>2</sub>, and shade (8), with lowest values observed in moist regions with dense canopy (17). Although concentration and δ<sup>13</sup>C values of atmospheric CO<sub>2</sub>

## Significance

**Humans evolved in response to the availability of plant and water resources over space and through time. Their influence on our species' evolution is debated, though, because archives of their spatial distribution are scarce at early human (hominin) localities. Meter-scale vegetation patterns are revealed from sedimentary plant biomarkers across an archaeological horizon at Olduvai Gorge (FLK Zinj). Biomarkers evince a varied local landscape with a woodland patch near a small freshwater wetland, surrounded by an open grassland landscape. Biomarkers from the wetland indicate diverse edible plants near potable water. The coexistence of butchered large animal bones and hominin remains, including juveniles, within an isolated biomarker-delineated wooded microhabitat at FLK Zinj provide support for early provisioning behaviors by our ancestors.**

Author contributions: C.R.M., G.M.A., M.D.-R., and K.H.F. designed research; C.R.M. performed research; C.R.M. and K.H.F. analyzed data; and C.R.M., G.M.A., M.D.-R., and K.H.F. wrote the paper.

The authors declare no conflict of interest.

This article is a PNAS Direct Submission.

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This article contains supporting information online at [www.pnas.org/lookup/suppl/doi:10.1073/pnas.1507055113/-DCSupplemental](http://www.pnas.org/lookup/suppl/doi:10.1073/pnas.1507055113/-DCSupplemental).

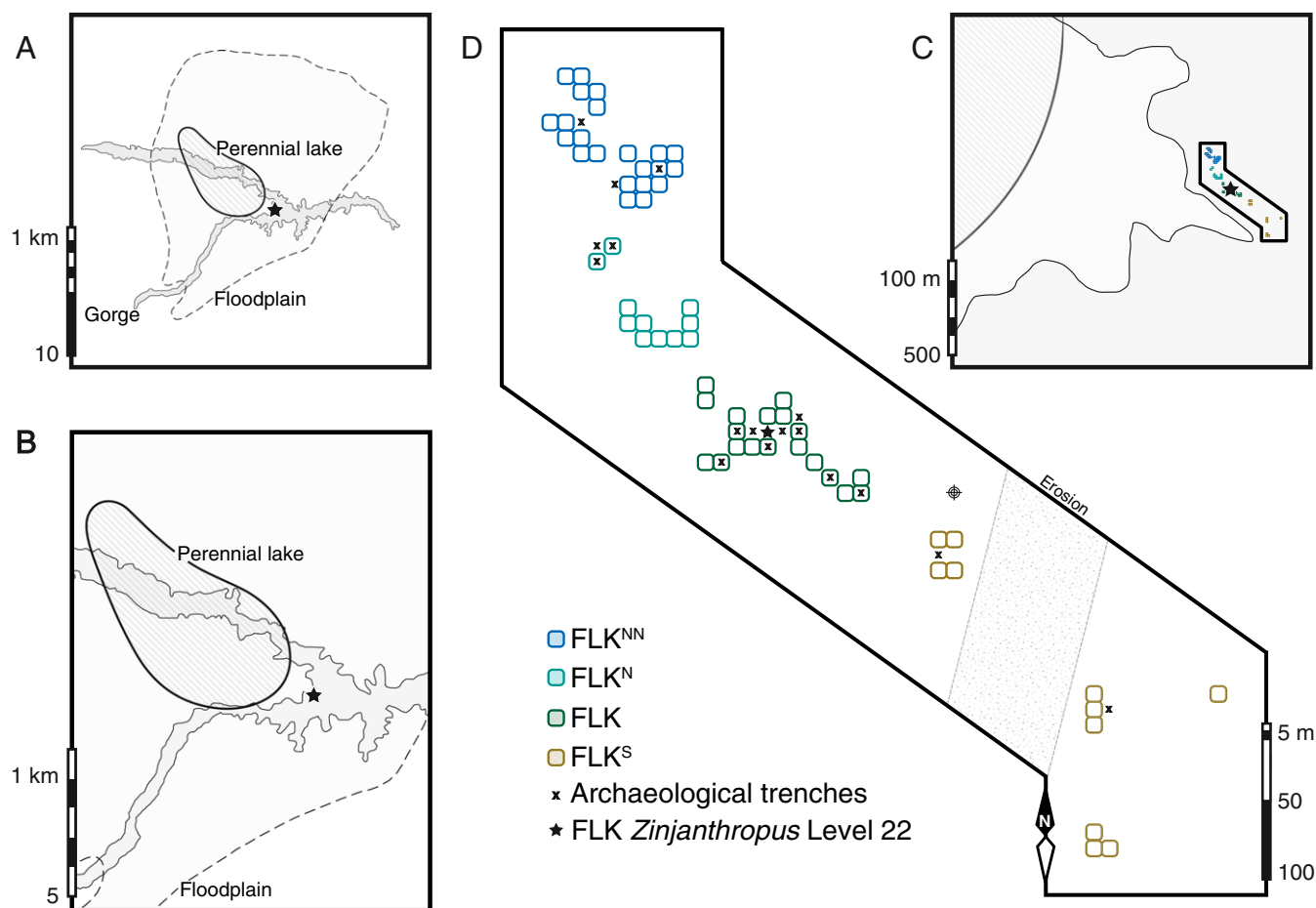
can affect  $C_3$  plant  $\delta^{13}C$  values (17), this influence is not relevant to our work here, which focuses on a single time window (*SI Discussion*). The large differences in leaf-wax  $\delta^{13}C$  values between closed  $C_3$  forest to open  $C_4$  grassland are consistent with soil organic carbon isotope gradients across canopy-shaded ground surfaces (6) and serve as a quantitative proxy for woody cover ( $f_{\text{woody}}$ ) in savannas (8).

As is observed for nonhuman primates, hominin dietary choices were likely shaped by ecosystem characteristics over habitat scales of 1–1,000 m<sup>2</sup> (3–5). To evaluate plant distributions at this small spatial scale (9), we excavated 71 paleosol samples from close-correlated trenches across a ~25,000-m<sup>2</sup> area that included FLK *Zinjanthropus* Level 22 (FLK *Zinj*) at Olduvai Gorge (Fig. 1). Recent excavations (18–21) at multiple trenches at four sites (FLK<sup>NN</sup>, FLK<sup>N</sup>, FLK, and FLK<sup>S</sup>, Fig. 1D) exposed a traceable thin (5–50 cm), waxy green to olive-brown clay horizon developed by pedogenic alterations of playa lake margin alluvium (22). Weak stratification and irregular redox stains suggest initial soil development occurred during playa lake regression (18, 22), around 1.848 Ma (ref. 23 and *SI Discussion*). To date, craniodental remains from at least three hominin individuals (18–20), including pre-adolescent early *Homo* and *Paranthropus boisei*, were recovered from FLK *Zinj*. Fossils and artifacts embedded in the paleosol horizon often protrude into an overlying airfall tuff (18, 19), which suggests fossil remains were catastrophically buried in situ under

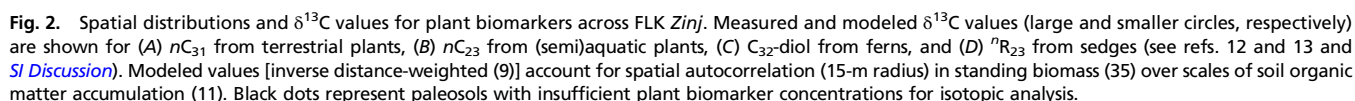
volcanic ash. Rapid burial likely fostered the exceptional preservation of both macrofossils (10) and plant biomarkers across the FLK *Zinj* land surface.

Plant biomarker signatures reveal distinct types of vegetation juxtaposed across the FLK *Zinj* land surface (Figs. 2–4 and Fig. S1). In the northwest, FLK<sup>NN</sup> trenches show high  $nC_{23}$   $\delta^{13}C$  values (Fig. 2B) as well as high  $C/V$  and  $P_{\text{aq}}$  values (Figs. 3 and 4A). They indicate floating or submerged aquatic plants (macrophytes) in standing freshwater (13), a finding that is consistent with nearby low-temperature freshwater carbonates (tufa), interpreted to be deposited from spring waters (22). Adjacent FLK<sup>N</sup> trenches have lower  $P_{\text{aq}}$  values (Fig. 4A) with occurrences of fern-derived  $C_{32}$ -diol and sedge-derived  $^{23}R_{23}$  (Fig. 2C and D). These biomarker distributions indicate an abrupt (around 10 m) transition from aquatic to wetland vegetation. Less than 100 m away (Fig. 1C), low  $nC_{31}$   $\delta^{13}C$  values (Fig. 2A) and low  $C/V$  and very low  $P_{\text{aq}}$  values (Figs. 3 and 4A) collectively indicate dense woody cover (Fig. 4B). In the farthest southeastern (FLK<sup>S</sup>) trenches, high  $C/V$  values and high  $\delta^{13}C$  values for  $C$  lignin phenols (Fig. 3) indicate open  $C_4$  grassland.

Biomarkers define a heterogeneous landscape at Olduvai and suggest an influence of local resources on hominin diets and behavior. It is recognized (2, 24–26) that early *Homo* species and *P. boisei* had similar physiological characteristics. These similarities in physical attributes suggest behavioral differences were what allowed for overlapping ranges and local coexistence



**Fig. 1.** Location and map of FLK *Zinj* paleosol excavations. (A and B) Location of FLK *Zinj* as referenced to reconstructed depositional environments at Olduvai Gorge during the early Pleistocene (18, 22) and the modern gorge walls. The perennial lake contained shallow saline-alkaline waters that frequently flooded the surrounding playa margin (i.e., floodplain) flats. (C) Outline of FLK *Zinj* paleosol excavation sites used for our spatial biomarker reconstructions. (D) Concentric (5 m) gridded distribution map of FLK *Zinj* paleosol excavations relative to previous archaeological trenches (18–21). Major aggregate complexes (FLK<sup>NN</sup>, FLK<sup>N</sup>, FLK, and FLK<sup>S</sup>) are color-coded to show excavation-site associations.



Alternative foodstuffs with abrasive,  $^{13}\text{C}$ -enriched biomass include seedless vascular plants (cryptogams), such as ferns and lycophytes [e.g., quillworts (27, 30)]. Ferns are widely distributed throughout eastern Africa in moist and shaded microhabitats (31) and are often found near dependable sources of drinking water (32). Today, ferns serve as a dietary resource for humans and nonhuman primates alike (27), and fiddlehead consumption is consistent with the inferred digestive physiology [salivary proteins (33)] and the microwear on molars (34) of

A scatter plot showing the relationship between S/V lignin phenols (y-axis, 0 to 1.5) and C/V lignin phenols (x-axis, 0 to 0.75). The plot includes data points for four lignin types: FLK<sup>NN</sup> (filled circles with a dot), FLK<sup>N</sup> (open circles), FLK (filled circles), and FLK<sup>S</sup> (circles with a cross). A color scale on the right indicates the  $\delta^{13}\text{C}$  value for p-Coumaric acid (C lignin phenol), ranging from -14 (yellow) to -20 (dark green). Four regions are labeled: W (top left), H (center), G (top right), and N (bottom left). Error bars are present for all data points.

**Fig. 3.** Molecular and isotopic signatures for lignin phenols across FLK *Zinj*. Bivariate plots are shown for diagnostic lignin compositional parameters (see refs. 12 and 15 and Table S1) associated with aggregate excavation complexes (Fig. 1C). Symbols are colored according to respective  $\delta^{13}\text{C}$  values for the C lignin phenol, *p*-coumaric acid. FLK symbols are uncolored due to insufficient *p*-coumaric acid concentrations for isotopic analysis. Representative lignin compositional parameters (12, 15) are shown for monocotyledonous herbaceous tissues (G), dicotyledonous herbaceous tissues (H), cryptogams (N), and dicotyledonous woody tissues (W).





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# Supporting Information

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## SI Materials and Methods

**Plant Biomarker Extraction.** Freeze-dried and powdered paleosols (~20–40 g) were extracted by accelerated solvent extraction (Dionex ASE 200 system) with DCM and methanol (90:10 by volume) in three cycles of 5 min at 100 °C and 1,500 psi (10.3 MPa). Then, extracted paleosols were oxidized under alkaline conditions with cupric oxide at 155 °C for 180 min in stainless steel pressure vessels (36) and then acidified with hydrochloric acid. Oxidation products (lignin phenols) were recovered by extraction with diethylether (36).

**Lipid Biomarker Isolation.** Total lipid extracts were separated into apolar (hydrocarbon), intermediate-polarity, and polar fractions over activated silica gel by elution with hexane, DCM, and methanol, respectively. Hydrocarbons were separated further into saturated and unsaturated fractions over activated silver-impregnated (5% by weight) silica gel by elution with hexane and DCM, respectively (37). In turn, saturated hydrocarbons were separated further into unbranched (*n*-alkanes) and branched fractions via a zeolitic (5 Å) sieve.

**Molecular Characterization.** Plant biomarkers were characterized by GC-MS with a Hewlett-Packard 6890 series GC and Hewlett-Packard 5973 mass selective detector. Samples were injected in splitless mode onto a 60-m DB5 fused-silica column (0.32 mm × 0.25 μm) via a Hewlett-Packard 7683 series autosampler. GC temperature was programmed to 60 °C for 1 min then ramped to 320 °C at 6 °C·min<sup>-1</sup> and held at final temperature for 20 min. Injector and detector temperatures were held at 320 °C. Functionalized biomarkers were derivatized with BSTFA and detected as trimethylsilyl (TMS) derivatives, and 5-*n*-alkylresorcinols were identified based on characteristic mass spectral fragment ions at *m/z* 268 (base peak) and *m/z* 281 (38). Midchain (1,ω20) diols were identified based on *m/z* 369 (base peak) and *m/z* 359 (39, 40).

**Isotopic Characterization.** Isotopic signatures were characterized by gas chromatography-combustion-isotope-ratio monitoring mass spectrometry with a Varian 3400 model GC connected to a Thermo MAT 252. Samples were injected in splitless mode onto a 60-m DB5 fused-silica column (0.32 mm × 0.25 μm) before combustion over nickel and platinum wire with oxygen in helium at 1,000 °C. Values were determined relative to reference gas calibrated to VPDB and expressed in permil (‰) units:

$$\delta^{13}\text{C} = 1,000(R_{\text{sample}}/R_{\text{standard}} - 1), R = {}^{13}\text{C}/{}^{12}\text{C}.$$

Within-run precision (1σ) and accuracy were determined by co-injected internal standards and are equal, respectively, to 0.14‰ and 0.09‰ (*n*-alkanes, *n* = 136), to 0.49‰ and 0.17‰ (diols; *n* = 48), and to 0.66‰ and 0.21‰ (lignin phenols; *n* = 40). Lignin phenol δ<sup>13</sup>C values are reported for vanillic, syringic, and *p*-coumaric acids (Table S1). Alcoholic and acidic functional groups of lignin phenols were derivatized using BSTFA. Isotopic corrections for carbons added by BSTFA were made via measurements of the δ<sup>13</sup>C of benzene-1,2-dicarboxylic acid [commonly called phthalic acid (Schimmelmann standards)] after derivatization and then correcting for the mass-balanced δ<sup>13</sup>C value of derivative carbon (15).

**Spatial Analysis.** Paleosol excavations were referenced to concentric (5 m) quadrats after geometric correction (0.42°) for tectonic deformation and sediment compaction (21). This approach

accounts for uneven spatial distributions in point density (41), enabling accurate distance representations in 2D interpolations (9).

Because plant biomarker reconstructions represent the sum of complex biogeochemical processes across space (9), we use a data-based (deterministic) spatial model based on inverse distance-weighted interpolations of all points within a fixed radius (see ref. 42). Based on combined theory (43) and modern observation (43–45), we calibrate our model with a 15-m radius to account for spatial autocorrelation in standing biomass over scales of soil organic matter accumulation (11). We converted spatially explicit point interpolations into probability surfaces assuming a unimodal (Gaussian) distribution (9, 41, 42). Derivative kernel density maps accurately represent primary spatial patterns (46) and facilitate comparison between diverse datasets with dissimilar resolution.

## SI Discussion

**Sedge Alkylresorcinols.** Resorcinolic compounds have been isolated from plants, algae, and bacteria (47). Odd-numbered long-chained 5-*n*-alkylresorcinol homologs are characteristic of the monocotyledonous plant (sub)families *Pooideae* and *Cyperaceae*. In eastern Africa, extant genera of *Pooideae* are limited to high-altitude environments [ $>2,000$  m (48)], well above the mean elevation of Olduvai Gorge (*ca.* 1,450 m). In contrast, extant genera of *Cyperaceae* are quite cosmopolitan, occurring in wetlands and seasonal floodplains from sea level to about 2,000 m (48, 49). Tropical Africa contains around 1,400 *Cyperaceae* species, of which ~50% occur in just two genera: *Cyperus* and *Carex* (50–54). Recent DNA data suggest that *C<sub>4</sub> Cyperus* form a monophyletic clade nested within *Cyperus sensu stricto* (53, 54) and comprise nine segregate genera. Together with earlier descriptive studies about the resorcinolic compounds of *C<sub>4</sub> Cyperus* sedges (e.g., *Cyperus longus*) in northern and East Africa (55, 56), our new analyses of *Cyperus papyrus* (Fig. S2) suggest that 5-*n*-alkylresorcinols with saturated C<sub>19</sub>–C<sub>25</sub> side chains occur in *Cyperus* sections of *Rotundi* Clarke and *Papyrus* (54). Earlier studies also indicate alkylresorcinols occur in taxa of the *Cyperus* segregate, *Pycreus* sedges (51, 55, 57) across northeast Africa (58). Importantly, <sup>23</sup>R<sub>23</sub> δ<sup>13</sup>C values for *C. papyrus* are much higher [−21.5 ± 0.5‰ (*n* = 3)] than measured at FLK Zinj (Fig. S2).

*Cyperus* and *Pycreus* (54) are within the cardinal tribe Cyperaeae, which is one of two major clades of the subfamily Cyperoideae (Fig. S3). The other major clade of this subfamily possesses cardinal tribes called Cariceae and Scirpeae (50, 53), which include widespread African sedges of the genera *Carex* and *Eriophorum* (58), respectively. We found *Carex trisperma* contains significant <sup>23</sup>R<sub>23</sub> with low δ<sup>13</sup>C values [−29.6 ± 1.1‰ (*n* = 4)], and this corroborates earlier data for cosmopolitan species of *Carex* and *Eriophorum* (38, 59, 60). Thus, 5-*n*-alkylresorcinols with saturated C<sub>19</sub>–C<sub>25</sub> side chains occur in (sub)tropical tribes of Cyperoideae that include *C<sub>4</sub>* and *C<sub>3</sub>* clades (53, 54).

**Fern Alkyldiols.** Midchain diols derive from a variety of marine, freshwater, and terrestrial photosynthetic organisms (40, 61) and can preserve in sediments and soils over geologic timescales (39). In general, longer midchain diols occur as a series of even-numbered homologs with characteristic isomer distributions (40). Previous studies identify terminal C<sub>30</sub>–C<sub>32</sub> diols with ω20 hydroxyl positions (e.g., C<sub>32</sub>-diol) as diagnostic molecular constituents in ferns (39, 61), and we identified the same major homologs in paleosols around FLK Zinj (Fig. S2).

Strict  $C_4$  photosynthesis is not found in monilophytes [ferns (62)] or related lycophytes (63), but these common vascular cryptogams (64) often use alternative carbon-concentrating pathways [e.g., crassulacean acid metabolism (CAM)] that lead to similarly  $^{13}\text{C}$ -enriched lipids and biomass (30, 65–70). In subtropical Africa, ferns can show variable degrees of CAM (69–71), and their associated bulk  $\delta^{13}\text{C}$  values span at least 14‰, from  $C_3$  ( $\sim -34$ ‰) to more  $C_4$ -like ( $\sim -20$ ‰) signatures (39, 71–73). Other edible vascular cryptogams with CAM [e.g., *Isoetes* sp. (14, 65)] likewise are common in seasonal freshwater semiaquatic environments across southern and East Africa (64, 74) and have even more  $^{13}\text{C}$ -enriched  $C_4$ -like signatures (30, 65, 70, 73).

**Sedimentation and Landscape Relief During Paleosol Formation.** Recent landscape reconstructions disagree on interpreted sediment facies and stratigraphic correlations around FLK Zinj (see refs. 18–21). These differences largely reflect different interpretations of the origin of local topographic relief. Here, we adopt the approach of Uribe-Larrea et al. (21), which links sediment facies analyses with geomorphic models (Fig. S4) for stratigraphic correlations and landscape reconstruction. An alternative model that FLK Zinj was developed by gradational fluvial incision or surface erosion (20) is not supported by this approach. Instead, cumulative evidence suggests low-energy processes dominated during sediment deposition at FLK Zinj and that a paleosol surface developed during lake regression from tuffaceous clays and was catastrophically overlain with volcanic ash [Tuff I<sup>C</sup> (21)].

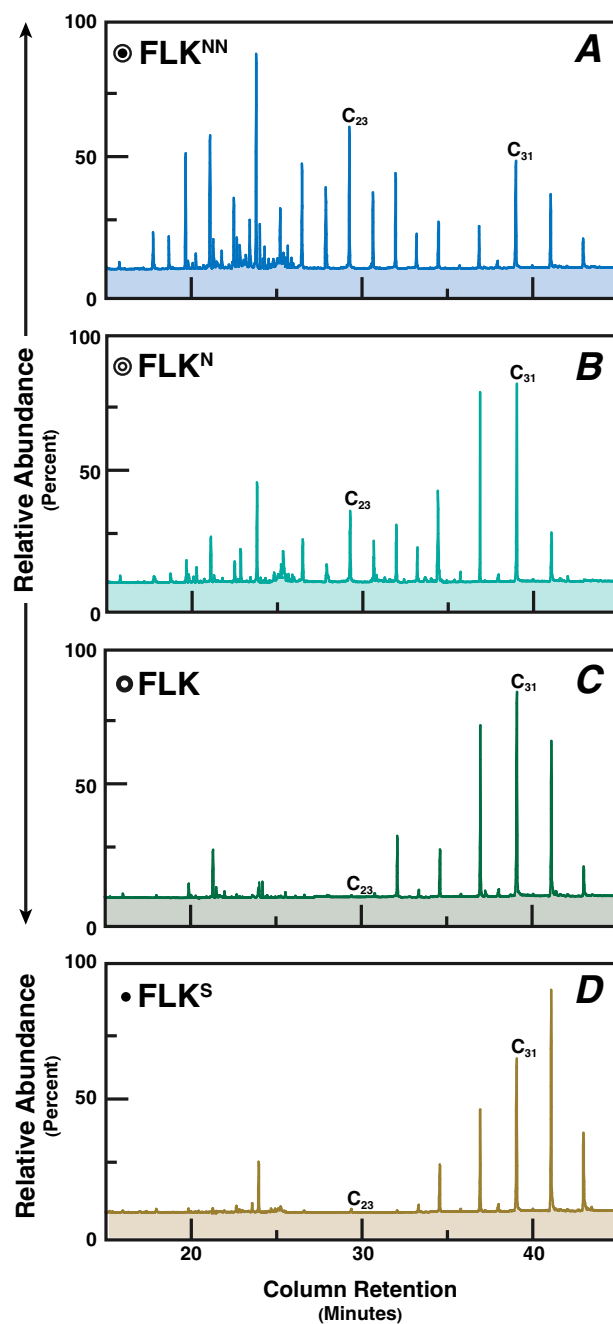
The exact timespan represented by paleosols around FLK Zinj is debated but constrained by high-precision radiometric ages and taphonomic data. Although Tuff I<sup>C</sup> lies about 100 cm above Tuff I<sup>B</sup>, they have indistinguishable argon–argon ages of  $1.848 \pm 0.008$  Ma and  $1.848 \pm 0.003$  Ma (23), which set the maximum duration of this entire sequence at 0.011 Ma (75). Taphonomic studies indicate archaeological materials at FLK Zinj accumulated during a time window with a duration that was somewhere between several successive seasons up to a century (4, 18–20, 76, 77). These short durations are estimated from an accumulation span of

50–100 y derived from the sedimentation rate ( $0.1\text{--}0.4\text{ mm}\cdot\text{y}^{-1}$ ) during paleosol formation at FLK Zinj (75, 78), the *ca.* 50-y residence time of organic carbon in modern (sub)tropic surface soil horizons [0–10 cm (6, 79–81)], and our average sample depth resolution ( $\sim 0\text{--}20$  mm underneath Tuff I<sup>C</sup>).

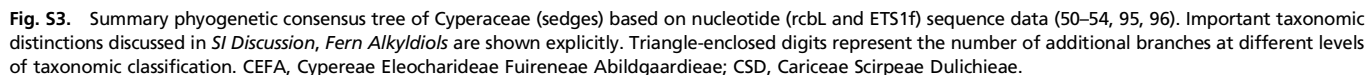
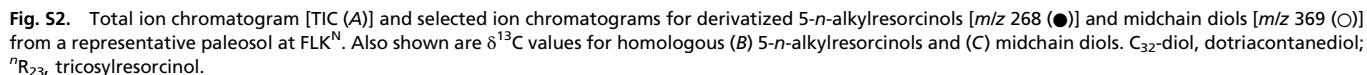
**Prior Constraints on *P. boisei* Dietary Behavior.** Nonspecific plant and animal fossils complement plant biomarker signatures to provide heuristic constraints on the local distributions and availability of food resources around FLK Zinj. Scattered leaf impressions bearing parallel longitudinal striations occur at southern FLK<sup>NN</sup> and FLK<sup>N</sup> (19, 20, 82) alongside *Heterocephalus* [African mole-rat (83)] fossils (18, 76) and corroborate plant biomarker signatures in suggesting that  $C_3$  graminoids [e.g., *Typha* (19, 82)] were locally abundant around FLK Zinj (Fig. 2 and Fig. S1). Further, silicified culms in sediments underlying FLK Zinj (17, 82) resemble the edible USOs (e.g., corms) of African cryptogams [e.g., *Isoetes* (84, 85)], which often show lower  $\delta^{13}\text{C}$  values due to  $C_3$ – $C_4$  intermediate photosynthesis (30, 65, 69, 70). Proximate playa lakeshore vegetation was an unlikely source for edible  $C_4$  grasses and sedges, because saline-alkaline water bodies [e.g., paleolake Olduvai (18)] are characteristically devoid of aquatic and emergent  $C_4$  macrophytes (86).

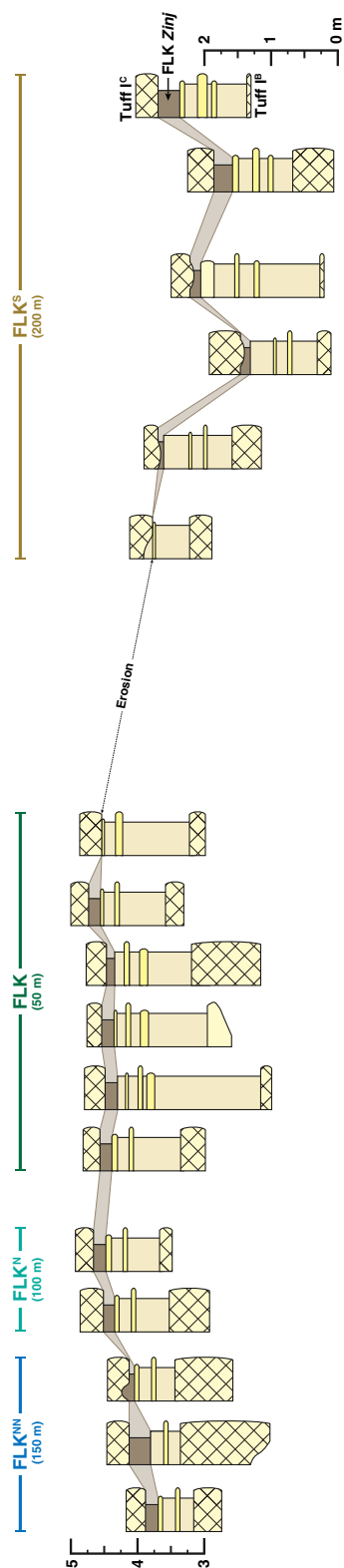
Dietary behavior in *P. boisei* is further informed by enamel  $\delta^{13}\text{C}$  values for contemporaneous animals in eastern and southern Africa during the early Pleistocene. Enamel  $\delta^{13}\text{C}$  values of *P. boisei* are markedly high relative to modern and fossil *Heterocephalus* species with specialized USO-rich diets (87), suggesting limited USO consumption by *P. boisei*. Besides, most wild USOs are innutritious without extensive extraoral processing or roasting (88). Rather, *P. boisei* shows enamel  $\delta^{13}\text{C}$  values intermediate between fossil *Hippopotamus* (hippos) and *Phaocochoerus* (pigs) in eastern Africa (24, 89). Based on combined isotopic (89) and craniodental (24) evidence from individual fossils at Olduvai Gorge, both taxa consumed tough browse-graze vegetation (90) such as ferns (91–94), suggesting similar dietary and characteristics for *P. boisei*.



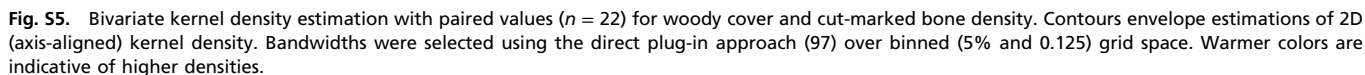


**Fig. S1.** Total ion chromatograms for saturated hydrocarbons in representative paleosols at (A) FLK<sup>NN</sup>, (B) FLK<sup>N</sup>, (C) FLK, and (D) FLK<sup>S</sup>. C<sub>23</sub>, tricosane; C<sub>25</sub>, pentacosane; C<sub>29</sub>, nonacosane; C<sub>31</sub>, hentriacontane.





**Fig. S4.** Schematic representation of the associated Tuff<sup>18</sup> to Tuff<sup>15</sup> stratigraphic sequences at FLK Zinj paleosol excavations after geometric correction (0.42°) for tectonic deformation and sediment compaction (2/1). The FLK Zinj soil horizon is shown in brown and traced between excavation sites. Excavation complexes are color-coded to match Figs. 1, 2, and 4). Distances shown under each excavation complex represent their associated (horizontal) transect lengths.



Plant biomarker	Source vegetation	Complex	$\delta^{13}\text{C}_{\text{Biomarker}}$			[Biomarker], $\mu\text{g}\cdot\text{g}^{-1}$ C
			Average	SD	$n$	
$n\text{C}_{23}$	Aquatic plants	FLK <sup>NN</sup>	-23.4	0.6	27	4
		FLK <sup>N</sup>	-24.7	1.2	12	3
$n\text{C}_{31}$	Terrestrial plants	FLK <sup>NN</sup>	-26.6	1.3	27	4
		FLK <sup>N</sup>	-29.7	1.1	12	4
		FLK	-33.8	0.9	19	7
		FLK <sup>S</sup>	-21.9	0.6	12	12
		FLK <sup>NN</sup>	-33.8	0.3	2	1
$^n\text{R}_{23}$	Sedges	FLK <sup>N</sup>	-32.0	0.9	5	2
		FLK <sup>NN</sup>	-26.7	0.5	5	2
C <sub>32</sub> -diol	Ferns	FLK <sup>N</sup>	-27.1	0.5	8	6
		FLK	-30.3	0.7	4	2
		FLK <sup>NN</sup>	-24.4*	1.0	4	182
V	Vascular plants	FLK <sup>N</sup>	-24.0*	—	1	207
		FLK	-33.6*	0.2	3	111
		FLK <sup>S</sup>	-17.5*	0.4	2	159
		FLK <sup>NN</sup>	-23.0 <sup>†</sup>	1.2	4	184
		FLK <sup>N</sup>	-24.9 <sup>†</sup>	—	1	93
S	Angiosperms	FLK	-33.2 <sup>†</sup>	0.2	3	84
		FLK <sup>S</sup>	-14.5 <sup>†</sup>	0.2	2	167
		FLK <sup>NN</sup>	-17.9 <sup>‡</sup>	1.3	4	55
		FLK <sup>N</sup>	-19.4 <sup>‡</sup>	—	1	47
		FLK <sup>S</sup>	-14.2 <sup>‡</sup>	0.1	2	92
C	Herbaceous					

<sup>‡</sup>*p*-coumaric acid (TMS derivative).